

STATE OF WYOMING
THE GEOLOGICAL SURVEY OF WYOMING
S. H. KNIGHT, State Geologist

BULLETIN No. 30

UNDERGROUND WATER
RESOURCES
— OF —
HORSE CREEK
— AND —
BEAR CREEK VALLEYS,
Southeastern Wyoming

BY
W. LYLE DOCKERY



UNIVERSITY OF WYOMING
LARAMIE, WYOMING
MAY, 1940

CONTENTS

	PAGE
Introduction	3
Topography and Drainage	4
Stratigraphy	5
Upper Cretaceous	5
Colorado group	6
Pierre shale	6
Fox Hills sandstone	6
Lance formation	6
Tertiary	6
General statement	7
Chadron formation	7
Brule formation	7
Arikaree formation	8
Age relations	9
Quaternary	9
Alluvium	9
Structure	10
Precipitation	10
Distribution	10
Disposal	13
Surface Water	13
Creeks	14
Springs	15
Records of stream discharge	16
Underground Water	16
Source	17
Relation to surface features	17
Occurrence	17
Colorado group, Pierre shale, and Fox Hills sandstone	18
Lance formation	18
Chadron formation	18
Brule formation	19
Brule-Arikaree contact	19
Arikaree formation	19
Alluvium	20
Depth to water table	20
Artesian conditions	20
Recharge from stream and irrigation water	21
Irrigation supplies	21
Municipal supplies	22
References	23
Well Data	23
Data for wells shown on Pl. II	27
Data for wells shown on Pl. III	27

ILLUSTRATIONS

	PAGE
Figure 1. Conglomeratic phase of the Arikaree near Underwood, sec. 31, T. 19 N., R. 69 W.	8
2. Cross-bedded sandstone phase of the Arikaree in Bear Creek Valley, sec. 4, T. 19 N., R. 63 W.	9
3. Spring (No. 89) issuing from the contact between a massive sandstone and the overlying thin-bedded fractured sandstone in the Arikaree, sec. 17, T. 19 N., R. 65 W.	15
Plate I. Map of Horse and Bear Creek region, southeastern Wyoming, showing hydrography and depth to water table	33
Plate II. Map of upper Horse and Bear Creek region, southeastern Wyoming, showing geology, hydrography, and location of wells Nos. 1-131	35
Plate III. Map of lower Horse and Bear Creek region, southeastern Wyoming, showing geology, hydrography, and location of wells Nos. 132-288	37

UNDERGROUND WATER RESOURCES
— OF —
HORSE CREEK
— AND —
BEAR CREEK VALLEYS,
Southeastern Wyoming*

By W. LYLE DOCKERY

INTRODUCTION

This report treats with the underground water resources of 1,355 square miles in Goshen, Laramie, and Platte Counties, Wyoming, and constitutes part of a program of investigation of the underground water resources of the state by the Geological Survey of Wyoming. The area comprises the drainage basin of Horse Creek and tributaries (Pl. I), excluding the part in T. 23 N., R. 62 W., northwest of the Fort Laramie Canal, and the western part of the Horse Creek drainage basin in the Laramie Range. There is also included the northern part of the Lodgepole drainage basin, which was not covered in a previous ground-water report on Lodgepole Valley (3).†

Horse Creek rises in the Laramie Range and flows eastward and northeastward across the Great Plains to the North Platte River near Morrill, Nebraska. Its length, not including meanders, is about 125 miles, of which approximately 100 miles is in the area here described. Bear Creek, the principal tributary of Horse Creek, rises in the Great Plains, flows eastward and joins Horse Creek near Lagrange, Wyoming. Its approximate length is 55 miles.

During the last decade, because of the success of irrigation utilizing ground water in Lodgepole Valley to the south and in the North Platte Valley in Nebraska to the east, there has been considerable interest and speculation about applying the same methods locally. The residents usually have a fairly accurate knowledge of well yields and some knowledge of the general depth to water. They are entirely uninformed, however, as to the geologic factors which control the occurrence and quantities of underground water. Consequently, many thousands of dollars have been spent for drilling wells in localities which proved to be unproductive of large-yield wells. It was largely with the view of attempting to alleviate this condition that the present study was undertaken. Field investigation during the summer of 1939

* Thesis submitted to the Department of Geology and the Committee on Graduate Study at the University of Wyoming, in partial fulfillment of requirements for the degree of Master of Arts.

† Numbers in parentheses refer to the list of references on page 22. A number following a colon in parentheses gives the page in the reference cited.

consisted of study of the geology with reference to water-supply problems, geologic mapping, and collection of data on 288 wells.

The region has a semi-arid climate typical of the higher Great Plains. The only weather station in the area is at Lagrange, elevation 4,500 feet. Here the mean annual temperature is 46°; the average dates of latest and earliest frosts are May 22 and September 25 respectively; the average length of the growing season is 126 days. Mean annual precipitation is about 17 inches; on an average there are 78 days in the year with 0.01 inch or more precipitation. Altitude variation from 4,060 to 7,000 feet in the surrounding region undoubtedly causes considerable differences in temperature, and minor differences in precipitation from the Lagrange figures.

The early settlers were ranchers who had access to the free range. Land was plentiful, the country thinly populated, and titles were obtained only for land needed for home ranches. These were generally in favored places close to the principal streams. Deep wells were later drilled on the plains some distance from streams and water for domestic and stock use was pumped into tanks by windmills or engines.

With the exception of the part of Horse Creek Valley below Meriden and several other areas where irrigation is possible, the land is best adapted to grazing.

The only available water-supply bulletin, which covers the eastern part of the area, is that by Adams (1) published in 1902. This work is chiefly a resume of the general geology and surface waters.

The writer wishes to thank Dr. S. H. Knight, State Geologist of Wyoming, for making the work possible, the several faculty members of the Geology Department, University of Wyoming, for valuable suggestions, especially Dr. R. H. Beckwith, who revised and criticized the manuscript, and Dr. C. H. Falkenbach of the Frick Laboratory, American Museum of Natural History, for information on age relations of the White River group near the Laramie Range. Appreciation is also extended to the residents of the area, whose cooperation and kindness contributed to the success of the work.

TOPOGRAPHY AND DRAINAGE

The northward-trending Laramie Range, with an average elevation of 8,000 feet, lies to the west of the area. In the extreme western part, late Paleozoic and Mesozoic rocks (Pl. II) dip steeply to the east along the east flank of the Laramie Range and form foothills several miles wide. The rocks are largely shales; the more resistant sandstones form hogbacks trending parallel to the mountain front.

A well-defined subsequent valley from two to five miles wide separates the foothills of the Laramie Range from the Great Plains and extends approximately 25 miles to the south. The valley floor

is a gently rolling surface cut in the late Mesozoic and Tertiary sediments. Because of numerous streams which flow eastward across the valley and the absence of an integrated drainage system, it has not previously been named. The name Federal Valley is suggested from Federal post office near the center of the valley seven miles south of Horse Creek post office. As here designated, Federal Valley extends from Granite Canyon post office on the Union Pacific Railroad 25 miles south of Horse Creek post office to the drainage divide between Horse Creek and Chugwater Creek five miles north of Horse Creek post office.

Federal Valley is bounded on the east by a westward-facing escarpment about 500 feet high cut across nearly horizontal Tertiary sediments. From the crest of the escarpment the gently undulating Great Plains slope eastward from an elevation of about 7,000 feet to 5,200 feet near the Nebraska line and are incised by the valleys of Horse Creek and Bear Creek. The east and northeast parts of the area have undergone more dissection; stream valleys broaden and isolated remnants of the higher plains surface form tops of buttes.

The North and South Forks of Horse Creek join at Horse Creek post office. In the first five miles downstream from here the valley narrows to a steep-walled trench up to 200 feet deep. In the next 40 miles the walls recede in a number of places and the valley widens to a mile or two. From Horse Creek post office to Meriden the north wall is more persistent and bolder than the south. Below Meriden the valley trends northeast and widens.

Bear Creek rises on the Great Plains in the western part of the area, is perennial for two and one-half miles, intermittent and ephemeral for 15 miles, and perennial for the remaining 37 miles to its junction with Horse Creek near Lagrange. There is a definite relation between the flow of Bear Creek and the shape of the valley. The upper perennial stretch is in a shallow canyon incised into the surface of the Great Plains; the 15-mile intermittent and ephemeral part is in a broad shallow valley; the lower 37-mile perennial part is in a narrow, steep-sided canyon, which terminates a few miles above the junction with Horse Creek.

Below the junction the valley of Horse Creek is part of Goshen Hole, a large depression excavated to a depth of nearly 1,000 feet below the Great Plains surface. It is bounded on the east, west and south by conspicuous escarpments and opens to the northeast into the North Platte Valley. Horse Creek enters Goshen Hole through a gap three or four miles wide in the south escarpment rim. Bear Mountain flanks the gap on the west and Sixty-six Mountain on the east.

STRATIGRAPHY UPPER CRETACEOUS

Colorado group.—The Colorado group, exposed in the vicinity of Horse Creek post office, includes the Thermopolis, Mowry, Frontier and Niobrara formations. As they are thin and could not

be shown individually on the map (Pl. II), they are shown as a unit. The Colorado group is about 950 feet thick and consists of dark carbonaceous shales, massive and slabby sandstones, and a few thin beds of limestone. Detailed sections of this and the two overlying Cretaceous units were measured by Loeffler (5). The sandstones and limestones carry small amounts of water.

Pierre shale.—The Pierre shale, which overlies the Colorado group conformably, consists of 4,000 feet of brown shales, sandy shales, and slabby sandstones. It dips eastward at about 50° and, despite its thickness, is of limited outcrop in the area under discussion. The formation carries little water and much of it is high in mineral matter. Water from the Harold C. Jones well (no. 2),* which produces from the Pierre, showed a standard calcium carbonate hardness of 720 parts per million and also Glauber's salt and Epsom salt. The water has approximately three and one-half times the hardness of that used for domestic purposes in Laramie.

Fox Hills sandstone.—The Fox Hills sandstone, a gray massive moderately hard sandstone containing numerous tan and dark-gray shale beds, overlies the Pierre conformably, has a thickness of 375 feet, and dips eastward at about 30°. It crops out only in the extreme western part of the area. The Fox Hills sandstone is known to yield water of better quality than that of the Pierre, but it is hard.

Lance formation.—The Lance formation, the youngest Cretaceous unit exposed, crops out only in lower Horse Creek Valley and in Goshen Hole. The formation consists of a succession of tan to yellow sandstones and shales. Lenticular oyster beds containing *Ostrea glabra*, a two-foot layer of impure lignitic coal, *Triceratops* remains, and cannonball concretions as large as two feet in diameter occur about five miles west of the Nebraska boundary. Schlaikjer (8: 51) says the Lance formation of Goshen Hole consists of a lower succession of continental beds 100 to 200 feet thick, brackish-water beds 80 to 125 feet thick, and an upper succession of continental beds 60 to 100 feet thick. The relation of the Lance to the underlying formations is not definitely known. The Lance formation is an aquifer, but the quality of the water varies from bed to bed; waters from some horizons are highly mineralized.

TERTIARY

General statement.—The Tertiary rocks are represented by the Chadron and Brule formations, of Oligocene age, and the Arikaree formation of late Oligocene or Miocene age. Because of lateral variations in lithology, gradational contacts, and poor exposures, the formations cannot be differentiated in many places.

* Well locations are shown on Plates II and III and locations by section, township, and range are given on pages 23 to 32.

The three were consequently mapped together as a single unit, the White River group.

Chadron formation.—The Chadron formation in the western part of the area near Horse Creek post office is a light-brown to drab locally cross-bedded poorly cemented arkosic conglomerate composed chiefly of material derived from the gigantic core of the Laramie Range. The thickness here is from 20 to 100 feet. To the east the Chadron is covered by younger rocks. It reappears around the edge of Goshen Hole, where it consists of lenses up to 60 feet thick of gray, greenish-gray, dark-red, and olive sandstones and sandy shales between the Lance formation and the Brule.

In the Goshen Hole region the Chadron contains an abundance of water of good quality.

Brule formation.—The Brule formation is typically a light-gray or light-tan massive moderately soft stratified andesitic and dacitic volcanic ash, which forms the lower slopes of the escarpments surrounding Goshen Hole and the escarpment near Horse Creek post office. It has an average thickness of about 200 feet and is more uniform in lithology than the other Tertiary formations. In many localities only a part of the formation is exposed and it is not always possible to tell which part is represented. At some places the Brule is entirely homogeneous; at others it contains lenticular hard layers or local sandy lenses. Near the Laramie Range the upper part interfingers with conglomerates similar in lithology to those of the overlying Arikaree formation.

The upper part of the Brule is cut by bed joints and at least one set of vertical joints. Recently weathered material consists of angular blocks up to a foot across that are bounded by joint surfaces. In addition there are fissures up to 12 inches wide that penetrate the upper part to depths of at least 30 feet. In some cases, they are filled by material from the overlying Arikaree, but in most, by material of the same lithology as the wall rock. The Brule is normally impervious. Local conglomerate lenses yield small amounts of water. Wells in strongly jointed and fissured areas yield large amounts of water. Water from the Brule is of good quality.

Arikaree formation.—The Arikaree formation overlies the Brule. In some places there is a disconformity between them; at others there is no evidence of an erosional break. The Arikaree has three textural phases. These are, from west to east, (1) coarse conglomerate (Fig. 1), (2) coarse evenly-bedded thin sandstones grading laterally into beds with large-scale cross-bedding, (Fig. 2), and (3) massive fine-grained gray salt-and-pepper sandstones with numerous resistant lenticular layers. The first phase is best developed in the area near Horse Creek post office, the second is well exposed in Bear Creek Canyon in T. 19 N., R. 63 W., and the third, which is eroded into pillars and spires, forms the summit of Bear Mountain west of Lagrange and is present farther east. These phases show only the general gradational aspects of the

formation. In addition vertical variations are found. The Arikaree forms the surface of the Great Plains and the upper sheer edges of the escarpments. Its thickness is variable, with an average of about 275 feet. Because of its high topographic position the formation yields little water, which is, however, of excellent quality.

Age relations.—The White River group along the east side of the Laramie Range is not thoroughly understood. Because of the disconformity separating the Arikaree and the Brule, and the limited faunal evidence, the age of the Arikaree was placed as Miocene. Recent detailed studies of the Tertiary sediments of Wyoming have shown that striking changes in lithology are not necessarily indicative of age difference. The disconformity separating the stream-deposited Arikaree and the subjacent Brule has no value as an indicator of a time break, for one would normally expect a disconformity beneath a stream-laid deposit of the magnitude of the Arikaree formation irrespective of previous subaerial erosion.

A study of the Tertiary deposits in southeastern Wyoming near the Laramie Range shows that the processes controlling the deposition of the White River group were complex and involved working and reworking of sediments by streams flowing from or across the mountains west of the region. Thus it is apparent that a fauna of a certain age found in a given lithologic type does not necessarily mean that the type is everywhere of the same age.

The most complete collection of vertebrate remains from the White River group of southeastern Wyoming was made by the Frick Laboratory, American Museum of Natural History, during



FIG. 1.—Conglomeratic phase of the Arikaree near Underwood, sec. 31, T. 19 N., R. 69 W.

the summer of 1939. Dr. C. H. Falkenbach, who directed the field work, expressed himself* as believing the White River group of southeastern Wyoming to be at least 90 per cent Oligocene in age.

The complex lithology, gradational contacts, intertonguing of sediments and scarcity of vertebrate remains lead the writer to believe that the terms Chadron, Brule, and Arikaree as applied to the Tertiary formations flanking the east side of the Laramie Range will never have more than general significance.

QUATERNARY

Alluvium.—The floors of Horse Creek and Bear Creek Valleys, except in canyons, are covered by alluvium and gravel. The thickness and character of the alluvium are variable. It was derived largely from the Tertiary rocks within the area, although it is probable that the Laramie Range has contributed to the silt and gravel along Horse Creek.

Adjacent to the stream channels the alluvium, from a few to 30 feet thick, contains water in considerable quantities. Water is released freely and is suitable for domestic and irrigation purposes.

STRUCTURE

The pre-Cambrian crystalline rocks in the core of the Laramie Range are flanked on the east by late Paleozoic and Mesozoic sediments dipping steeply to the east. These are unconformably overlain by almost horizontal Tertiary sediments of the White

* Oral communication, June, 1939.

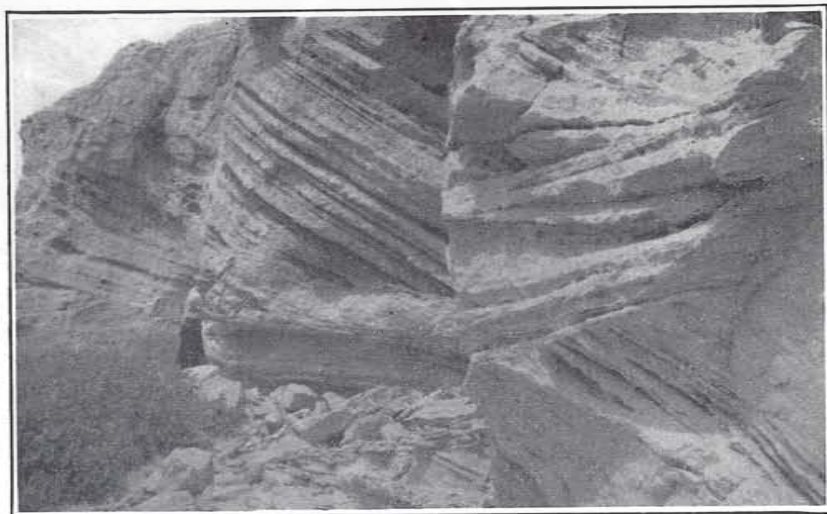


FIG. 2.—Cross-bedded sandstone phase of the Arikaree in Bear Creek Valley, sec. 4, T. 19 N., R. 63 W.

River group. In places the Tertiary sediments extend up against the mountains and completely cover the Paleozoic and Mesozoic rocks. The older rocks, because of the overlapping Tertiary, are of small surface areal extent.

The White River group in the westward-facing escarpment near Horse Creek post office dips eastward at an angle less than that of the Great Plains. In the eastern part of the area, in Goshen Hole, the Tertiary rocks have been removed by erosion, exposing the late Cretaceous Lance formation. The Lance is poorly exposed because of deep weathering and residual material from the Tertiary; little is known about the structure except that the prevailing dips are low.

The erosion surface upon which the Tertiary sediments were deposited was one of marked relief, 315 feet at an exposure one mile northeast of Farthing, which is near the Laramie Range about 10 miles north of Horse Creek post office. It is probable that farther east of the mountains the surface has less relief.

PRECIPITATION DISTRIBUTION

The only precipitation data available within the area are those for Lagrange given below.

There is a range from a minimum of 11.37 inches in 1911 to a maximum of 25.05 inches in 1912. It is of interest to note that the minimum and maximum occurred in successive years. The mean of the annual precipitations for the 18 scattered years of complete yearly records is 17.65 inches, which is comparable to the mean annual of other stations in neighboring sections of the Great Plains province. For nine years the precipitation was above the mean annual and for nine it was below. During three years there was more than 20 inches of precipitation; during five there was less than 15 inches.

Approximately 71% of the annual precipitation occurs between April and September inclusive. May, with a mean monthly precipitation of 2.51 inches is the wettest month; January, with 0.49 inch, is the driest. On the other hand, during May, 1930, there was only 0.20 inch, and during January, 1921, there were 1.58 inches. The largest precipitation recorded in a single month was 6.84 inches during May, 1917, and there are two instances of more than 5.00 inches in August.

DISPOSAL

Precipitation is disposed of through evaporation, transpiration, run-off, and percolation, which vary from place to place and with the seasons.

As 71% of the precipitation occurs from April to September inclusive, when the optimum conditions for evaporation prevail, a greater percentage of the precipitation evaporates than would be

the case if the period of maximum precipitation were in winter. Growing plants take water from the soil and discharge it to the atmosphere through their leaves as transpiration, which varies in amount with different plants and with the same plant under different environmental conditions. Transpiration reaches a maximum during the growing season. The amount of water lost by evaporation and transpiration is difficult to estimate.

Precipitation for Lagrange, Goshen County, Wyoming
Elevation 4,720 feet
(From records of the U. S. Weather Bureau)

Year	Jan.	Feb.	Mar.	April	May	June	July	Aug.	Sept.	Oct.	Nov.	Dec.	Annual
1902					1.20	2.78	3.46	0.50	3.40	0.20		1.43	
1910					1.65	0.60	1.52	1.00	3.28	0.82	0.18	0.79	
1911	0.77	0.16	0.03	1.38	1.06	1.63	0.96	1.08	0.86	2.64	0.30	0.50	11.37
1912	1.07	0.97	1.62	2.02	2.40	2.02	2.56	5.31	4.30	1.75	0.52	0.51	25.05
1913	0.41	0.75	0.49	0.86	2.99	1.94	3.73	2.73	1.38	0.53	0.22	2.59	18.62
1914	0.08	0.40	0.41	3.66	0.96	3.10	1.49	1.00	4.42	1.83	0.20	0.30	13.85
1915	0.49	0.69	1.18	3.99	2.62	3.47	1.35	1.96	4.10	3.01	0.27	0.53	23.66
1916	0.19	0.31	0.30	0.85	2.96	0.25	2.54	2.35	0.45	1.28	0.57	0.33	12.38
1917	0.22	0.10	0.53	1.37	6.84	4.15	1.09	0.89	0.47	0.38	0.37	0.43	16.84
1918	0.71	0.35	0.44	2.72	2.68	3.39	1.92	1.23	3.90	1.07	0.40	1.08	19.89
1919	0.04	0.75	1.72	1.48	0.21	1.01	1.69	0.08	1.10	1.33		0.60	
1920	0.24		0.77	4.75	2.46					0.02	0.64	0.50	
1921	1.58	0.32	0.26	1.74	2.17	1.24	0.92	1.01	0.90	0.81	0.56	1.62	13.13
1922	0.72	0.74	0.49	2.42	3.58	1.77	1.89	1.33	0.22	0.47	2.68	1.16	17.47
1923	0.25	1.37	0.70	1.95	2.80	4.60	0.86	1.22	1.26	3.04	0.73	0.81	19.59
1924	0.41	0.60	1.39	1.46	2.24	0.93	0.38	0.34	3.90	2.59	0.44	1.12	15.80
1925	0.13	0.68	0.86	1.51	1.97	2.86	1.35	1.63	0.89	2.62	0.83	2.18	17.51
1926	0.92	0.68	0.45	1.89	2.62	4.44	1.45	0.82	0.87	2.61	1.22	0.60	18.57
1927	0.30	0.44	2.90	5.45	2.21	2.71	1.65	4.11	0.88	1.36	0.46	0.46	22.93
1928	0.44	1.00	0.75			3.63	2.39	0.69	0.38				
1929						1.31	0.84	1.00					
1930	0.85	0.31	0.20	1.30	4.68	0.59	0.88	5.22	1.70	3.12	0.50	0.05	19.40
1931							0.69	0.77	0.50	1.66	0.25		
1932												0.65	
1933	0.02	0.09	1.40										
1936									0.58	0.47	0.88	1.11	
1937	0.48	T	0.94	1.04	1.62	2.50	2.66	0.16	0.88	1.11	0.71	1.15	13.25
1938	0.30	0.19	0.93	2.81	5.00	1.06	1.77	1.00	3.56	0.15	1.09	0.61	18.47
1939	0.70	1.07	0.95	0.48	0.78	1.41	2.07	1.13	0.60				
Mean	0.49	0.57	0.86	2.15	2.51	2.22	1.68	1.54	1.63	1.45	0.64	0.88	17.65

Part of the water that falls in the area is carried away by Horse Creek. Comparison of precipitation records and stream discharge at Horse Creek gaging station shows that as small a part of the precipitation as 1.7% leaves the area as run-off. This figure is probably unreliable, as the gaging station is located below Hawk Springs reservoir, from which water is diverted for irrigation. Part of the diverted water undoubtedly returns to the stream below the station.

The amount of water that penetrates to the water table is controlled by a number of factors, such as relief, porosity and

permeability of rocks and soil, plant cover, and water table gradient. Theis (9: 564-568) has made estimates of the amount of ground-water recharge from precipitation in the southern High Plains of New Mexico, Texas, Oklahoma, Kansas, and Colorado. The studies deal principally with an area of 35,000 square miles in the Llano Estacado, western Texas, of which Theis says:

"The main ground-water body in the southern High Plains lies in the Ogallala formation, of Pliocene and perhaps Miocene age. . . . The formation is variable laterally, but consists of four types of sediments, namely: Gravel found in old stream courses throughout the section, but most prominent near the base in the pre-Ogallala valleys; unbedded and unconsolidated fine sand and silt, which make up the bulk of the formation; cemented zones of calcareous fine sand and silt, differing from the last only in cementation, and forming caliche zones; and finally at the top of the formation a limestone several feet thick. . . . Overlying this limestone in most of the area is a deposit of fine sand or silt, reaching a thickness of over 100 feet, which has been correlated with the Pleistocene loess in Kansas and is probably of the same age throughout the area. The water-bearing Ogallala deposits rest generally on impervious shales ranging in age from Permian to Cretaceous. The surface of these pre-Ogallala beds slopes east and east-southeast about ten feet to the mile.

"The High Plains are in most places surrounded by scarps a few hundred feet high reaching down to pre-Ogallala rocks. The water-body in the Ogallala is thus raised above and isolated from the ground water in contiguous areas. The surface of the Plains, and particularly the Llano Estacado, is largely unscoured by erosion. It slopes also east and east-southeast about ten feet to the mile. The surface is smooth and flat except for a few sand dunes and for numerous sinks in the surface—'dry lakes' or 'buffalo wallows'—from a foot or two to over a hundred feet deep and from a few feet to several miles in diameter.

"Nearly all the precipitation on this area, especially on the Llano Estacado, is either discharged by transpiration and evaporation within a short distance of where it fell or seeps to and recharges the ground-water body. Only close to the few narrow-valleyed streams does any water run off. A good part collects as ephemeral ponds in the depressions, where it usually, but not in all localities, remains through months of drouth. To penetrate to the ground water, the precipitation must pass through the tight loess cover and through the dense but jointed limestone capping the Ogallala. In a few areas sand dunes cover the surface, and in these places recharge is more favored. The average yearly precipitation ranges from 14 inches in the southwestern part of the area to 22 inches in the northeastern. It occurs preponderantly as summer showers

and hence at the most inopportune time for rainfall penetration.

"The water table in the central part of the High Plains lies about halfway between the surface of the ground and the base of the Ogallala. Like the ground surface and the floor of the water body, it slopes east or east-southeast about ten feet to the mile.

"Natural discharge of the ground water occurs to a small extent in a few water-table lakes in depressions in the High Plains, but largely it must occur as seeps and springs at the eastern escarpment of the Plains. However, there are very few springs of any size and little evidence in the character of the vegetation of a great amount of seepage."

Quantitative estimates of ground-water recharge were made in the Llano Estacado by (1) measuring the flow across a certain cross-section, and dividing this by the tributary area, and (2) by measuring water table fluctuations. The estimates of the annual ground-water recharge from precipitation as one-eighth to two-thirds inch, which is about $3\frac{1}{2}\%$ of the 14-inch mean annual precipitation.

Topography, general structure, prevailing sandy nature of rocks and soil, precipitation, and position of water table in the Horse Creek drainage basin are similar to the features described in the Llano Estacado. The limestone of the Ogallala and overlying mantle of loess found in the southern High Plains do not have complements in the Horse Creek area; ephemeral ponds are also absent. The Horse Creek area contains a number of small springs, and the slope of the catchment area is higher, about 25 feet to the mile; the effect of this steeper gradient may be offset, however, by the permeable nature of the exposed formations.

Assuming a recharge rate of $3\frac{1}{2}\%$ of the mean annual precipitation, there would be an annual ground-water recharge of about 45,000 acre feet in the Horse Creek area; if the recharge were 5% there would be approximately 65,000 acre-feet. It is probable that ground-water recharge in the Horse Creek drainage basin is near these estimates.

SURFACE WATER CREEKS

Horse Creek is supplied in part by tributaries rising in the Laramie Range and in part by tributaries rising in the Great Plains. It is perennial throughout its length. A dam below the junction with Bear Creek diverts water into Hawk Springs reservoir, which has a capacity of 19,400 acre-feet. Water from the reservoir is used for irrigation in the vicinity of Hawk Springs.

Bear Creek heads near the western escarpment of the Great Plains, is fed by small springs along its upper part, and has perennial, intermittent, and ephemeral parts above its junction with

the tributary in sec. 9, T. 19 N., R. 66 W. Below the junction it is perennial, and receives water from numerous springs along the canyon walls. Water is diverted for local irrigation along the perennial parts.

Fox Creek is the only relatively large stream rising in the area which is perennial throughout its length. This is caused by the large springs near its head and smaller springs lower down. Only a small amount of the water is used for irrigation, and consequently the creek contributes a proportionately large flow to Bear Creek.

SPRINGS

The springs along upper Horse Creek Valley are small. In most cases they are near the tops of the escarpments far back from the stream, and the water evaporates or is absorbed near the springs. In some instances ranchers have constructed small reservoirs or tanks for livestock use near the springs. The springs along Horse Creek between Meriden and Hawk Springs are larger; springs are of small importance below Hawk Springs.

Hawk Springs, now covered by the waters of Hawk Springs reservoir, was located three and one-half miles north of Lagrange. Adams (1: 27-28) wrote concerning it in 1902:

"Hawk Springs, at the west end of 66 Mountain, is perhaps the largest so-called spring in the Goshen Hole country. . . . The ground is seepy, and the water rises in a number of places. . . . The source of the water cannot be definitely determined . . . (but) it appears that a considerable quantity of ground water passes laterally from the creek (Horse Creek) into gravels and sands, and it is not improbable that Hawk Springs is the point of issuance of this water, which follows a natural channel now obscured by surficial material."

Contact springs (7: 51) along Bear Creek and Fox Creek occur at the contact between the Brule and Arikaree formations. The Arikaree sandstone and conglomerate is more permeable than the underlying Brule ash and clay, and, in the absence of fissuring in the upper part of the Brule, water issues from the ground at the contact. These springs are small and only moisten the clay at the contact. At a distance this contact appears as a line of vegetation along the faces of the escarpments.

The larger springs along Fox Creek and Bear Creek occur within the Arikaree formation, generally at the contact of thin-bedded and underlying massive sandstone. The spring (no. 89) at the Ira P. Trotter ranch is one of these. (Fig. 3)

The spring (no. 162) at the C. C. Donahue ranch on Little Horse Creek four miles southwest of Meriden is estimated by Mr. Donahue to have a flow of $\frac{3}{4}$ second-foot. The flow of the spring (no. 258) at the Jelmer Johnson ranch on Horse Creek six miles south of Lagrange is of about the same magnitude.

RECORDS OF STREAM DISCHARGE

At the time of writing, records are available on the discharge of Horse Creek for the years 1929 to 1936 inclusive (10). A gaging station was located six miles southeast of Yoder, Wyo., in sec. 13, T. 22 N., R. 62 W., and was in operation from May, 1928, until September, 1932. This station measured the discharge of a drainage area of 1,320 square miles. Discharge data from this station for the years 1929 to 1931 inclusive, follow:

HORSE CREEK NEAR YODER, WYOMING

Discharge in acre-feet

Month	1929	1930	1931
January	984	2,820	2,670
February	833	4,220	1,090
March	2,520	2,370	3,580
April	6,430	228	1,830
May	7,070	836	437
June	6,600	281	54
July	234	15	12
August	25	3,310	29
September	1,900	2,950	22
October	258	3,010	312
November	3,230	2,730	655
December	3,390	2,840	246
Total	33,474	25,610	10,937

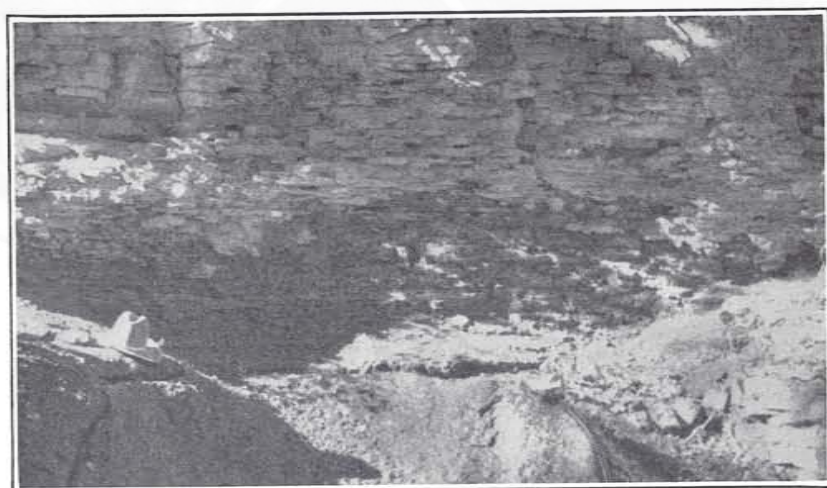


FIG. 3.—Spring (no. 89) issuing from the contact between a massive sandstone and the overlying thin-bedded fractured sandstone in the Arikaree. Sec. 17, T. 19 N., R. 65 W.

The other gaging station on Horse Creek is located in sec. 25, T. 23 N., R. 58 W., three miles northeast of Lyman, Neb., and measures the discharge of a drainage area of 1,860 square miles. Discharge data for the years 1932 to 1936 inclusive, follow:

HORSE CREEK NEAR LYMAN, NEBRASKA

Discharge in acre-feet

Month	1932	1933	1934	1935	1936
January	1,680	1,540	1,410	581	670
February	3,850	989	1,080	563	450
March	1,910	1,860	1,200	458	793
April	1,510	2,090	1,090	1,440	845
May	3,280	10,500	516	2,190	949
June	10,200	7,740	1,110	10,080	5,950
July	6,640	5,950	1,000	2,680	1,680
August	7,750	8,240	1,030	2,070	2,380
September	8,510	14,300	768	2,480	1,880
October	5,210	5,480	781	1,790	1,860
November	2,160	1,830	631	1,360	1,120
December	1,320	1,550	732	930	841
Total	54,020	62,069	11,348	26,622	19,418

Diversion of water for irrigation at Hawk Springs reservoir, 11 miles above the Yoder gaging station, influenced the records of the Yoder station. Gagings were abnormally low during some summer months when almost all the flow of Horse Creek was diverted into Hawk Springs reservoir. It is probable that return flow from irrigated areas added to the run-off during the fall months.

Records of the Lyman gaging station are more reliable than those of the Yoder station. Part of the water diverted at Hawk Springs reservoir re-enters Horse Creek as return flow above the Lyman station. The gagings at this station, however, undoubtedly include some water from outside the Horse Creek drainage area, as some of the land irrigated by the Fort Laramie canal, which diverts water from the North Platte River, provides a return flow into the lower part of Horse Creek.

UNDERGROUND WATER SOURCE

The underground water of the region is derived from two sources. Horse Creek contributes a relatively small but stable volume of water from the Laramie Range. Influent seepage from the stream is of importance in the valley near Horse Creek post office, and farther down where water is diverted for local irrigation.

The major source of underground water is, however, precipitation that falls on the area. Part of the precipitation seeps into the ground and percolates downward. Absorption is small in areas of dense clay and is proportionately high in outcrops of conglomerate and sandstone. For the entire area, an estimated ground-water recharge of 5% of the precipitation has been mentioned previously in this paper.

RELATION TO SURFACE FEATURES

Underground water is controlled to a large extent by surface features. Absorption is less on steep slopes than on gentle slopes, other factors being equal. Leakage from water-bearing beds occurs at springs in some places along escarpment faces and valley walls, especially along Fox Creek and Bear Creek. Sharp and narrow crests, which tend to produce well delineated drainage divides, are absent, but the long, narrow valleys form natural lines of drainage. Underground water normally moves toward them in much the same direction as surface water.

OCCURRENCE

Colorado group, Pierre shale, and Fox Hills sandstone.—These rocks, which consist of shales, sandstones, and a few thin beds of limestone, are exposed in the extreme western part of the area near Horse Creek post office (Pl. II), where they form foothills against the Laramie Range and are exposed in the floors of Federal Valley and upper Horse Creek Valley. They have a total thickness of about 5,300 feet and dip eastward at 30° to 60°; the dip decreases eastward from the Laramie Range.

There are no wells that produce from the Colorado group. The sandstones and limestones of the Colorado group are probably aquifers, but are at too great a depth under most of the area to be important sources of water. Wells drilled in the Pierre shale have had varying degrees of success, depending upon whether or not a sandstone aquifer was reached. In some places it is possible to locate wells so that they will produce from sandstones in the Pierre shale; at other places it is covered by alluvium or White River sediments and consequently the position of the sandstone beds cannot be determined from surface observation. Ground-water conditions in the Fox Hills sandstone are normally more favorable for well production than in the Pierre.

Tests on water from the Pierre show the presence of Epsom salt and Glauber's salt, and a standard CaCO_3 hardness of 720 parts per million, which is three and one-half times the hardness of that used for domestic purposes in Laramie. The Fox Hills yields water of better quality than the Pierre, but it is hard. Water from these formations is used for domestic purposes, but it leaves much to be desired. These rocks do not contain underground water in quantities sufficient for irrigation.

Lance formation.—The Lance formation, a succession of sandstones and sandy shales, is exposed in lower Horse Creek Valley and in Goshen Hole (Pl. III). It is the youngest Cretaceous unit exposed. The structure and stratigraphy of the Lance are not completely understood, but the prevailing dips are low and the average thickness is about 400 feet. It is doubtful that the entire succession is present at any one locality. Conditions for water production from shallow wells are in general more favorable than in the older Cretaceous rocks near Horse Creek post office. The quality of the water, which is not as good as that from the Tertiary formations, varies at different stratigraphic levels in the Lance. Water from some horizons is highly mineralized. There is abundant ground-water for domestic use, but it is doubtful that wells in the Lance would provide enough water for irrigation.

Chadron formation.—There is no record of wells drilled in the Chadron formation in the western part of the area. In the eastern part, around the edge of Goshen Hole, the Chadron occurs as lenses of sandstones and sandy shales up to 60 feet thick between the Lance and Brule. Wherever the Chadron extends below the water table, it carries water in quantities ample for domestic use. The water is of better quality than that of the underlying Cretaceous rocks.

Brule formation.—The average thickness of the Brule is about 200 feet. For the most part, it is a claystone composed largely of dacitic and andesitic volcanic ash, and is too impervious to permit free transmission of water. At some exposures, the Brule is cut by bed joints and at least one set of vertical joints; in addition, fissures up to 12 inches wide and 30 feet deep are present. In some instances, the fissures are filled with conglomerate and sandstone from the overlying Arikaree formation, but usually the filling is the same type of material as the wall rock.

It is believed that surface water percolates downward and enters the interconnecting joints, which function as narrow channels and transmit water to the wider fissures. The fissures are large enough to retain a considerable quantity of water and to concentrate the water supplied by the tributary joint system. A well which penetrates a fissure yields water; one that does not is dry or gives a small yield.

This hypothesis would explain the erratic success of wells drilled in the Brule, for often large-yield wells are located near dry holes. As the fissures are only a few inches wide, it seems improbable that large-yield wells produce only water stored in fissures. If such were the case, water in a fissure would behave as though stored in a tank, for the drainage area of the fissure walls is relatively small. On the other hand, if the fissure were connected by a tributary system of joints, the total drainage area of the joint walls would be much larger. The cone of influence of a well producing from the fissure would also be much larger.

The joints and fissures in the Brule are secondary, rather than original interstices, but their origin is not yet understood. The fissures are covered and their extent and size cannot be determined from surface observation and, consequently, the success of wells in the Brule cannot be anticipated in advance. Water from the Brule is of good quality, and is suitable for domestic use.

Brule-Arikaree contact.—Water percolating downward through the Arikaree may be held in the basal gravels or sandstones of the formation. Such a condition prevails where the upper part of the Brule is free from joints and fissures. Wells drilled in the Arikaree to the Arikaree-Brule contact usually yield enough water for domestic use, although the supply is commonly smaller than that obtained from fissures in the Brule. Some wells have failed to produce water in either the basal Arikaree or in the upper Brule. Such failures cannot be predicted.

Arikaree formation.—The Arikaree formation has an average thickness of about 275 feet, and forms the surface of the Great Plains. It is conglomeratic in the western part of the area (Fig. I). The conglomerate lenses pinch out eastward and are replaced by sands. In the eastern part of the area the conglomerates are represented by thin lenses of gravel. Here the formation is composed, for the most part, of soft permeable sandstones. It appears that both the conglomerate and sandstone facies are highly permeable, for on the Great Plains wells producing from the Arikaree formation are about 150 feet deep. This shows that the zone of saturation occurs normally in the lower part of the formation.

Local indurated layers in the Arikaree, such as those exposed at Bear Mountain and in adjacent areas, may impede or prevent downward percolation of ground water and produce perched water tables of small area. Perched water tables also occur where thin-bedded sandstones overlie more massive sandstones.

A few wells drilled in the Arikaree have encountered water under hydrostatic pressure, which caused the water to rise several feet in the wells; this condition, however, is not general. Because of the formation's high topographic position, it does not yield large amounts of water. The water is of good quality and occurs in quantities ample for domestic use, but insufficient for well irrigation.

Alluvium.—The streams, especially in the eastern part of the area, are aggrading, but have not yet developed wide flood plains. Deposits of alluvium along the streams, for the most part, are derived from rocks in the area; those of Horse Creek Valley are derived in part from the granites of the Laramie Range. The thickness of the alluvium ranges from a few feet to 30 feet. Twelve to 15 feet is the average thickness along the principal streams.

The alluvium is an important source of underground water in lower Horse Creek Valley near Lagrange. Here Horse Creek is in places an influent stream, i. e., contributes water to the zone of saturation. Water also enters the alluvium by seepage from the

adjacent uplands. Large amounts of gravel are present in the alluvium, so that rapid withdrawal of a large volume of water is possible. The two successful irrigation wells (nos. 251, 252) in the area produce largely from gravel beds in the alluvium.

Many residents secure water for domestic use from the alluvium. Well irrigation from the alluvium is most feasible near Lagrange. It is probable that irrigation wells might produce successfully from the stream alluvium elsewhere, provided the volume of gravel is large enough.

DEPTH TO WATER TABLE

The water table lies at depths ranging from a few feet to more than 200 feet beneath the surface. Plate I, compiled from well data, shows depth to the water table classified as under 50 feet, 50 to 100 feet, 100 to 200 feet, more than 200 feet, and unclassified. Areas designated as unclassified are those for which well data are lacking. In most instances, however, an approximate figure for depth to the water table can be obtained by reference to a near-by area for which the depth is known.

The shape of the water table normally conforms to that of the ground surface, but has less relief. The water table is nearer the surface in the valleys. Fluctuations of the water table are due to the different modes of occurrence of underground water just discussed, and because of the varying factor of water disposal through evaporation, transpiration, percolation, and run-off. Fluctuations could not be noted within the short time allotted to this study, and the residents have little knowledge of variation in water table levels. Mr. Hugh Stemler states that from observation of his irrigation wells (nos. 251, 252) in the alluvium of lower Horse Creek Valley, the water table is normally three feet lower in October than in May, provided precipitation is normal.

ARTESIAN CONDITIONS

The only flowing well in the area (no. 225) is in lower Horse Creek Valley about seven miles west of the Nebraska line. According to Mr. A. Jones, the owner, the well was drilled for oil and abandoned at a depth of about 5,000 feet. The flow is small, and the water is heavily mineralized. Water reaches the surface through a pipe two inches in diameter, and is used for watering livestock and for pasture irrigation. No log was available, and the horizon from which the water comes is not known.

The geology and topography are such that it is not anticipated that flowing artesian wells will be encountered in the Tertiary rocks. Deep wells sunk into the underlying Cretaceous rocks might flow, but it is probable that water from such wells would be unfit for domestic use.

RECHARGE FROM STREAM AND IRRIGATION WATER

Horse Creek and Bear Creek are influent in some places and effluent in others. A stream or stretch of a stream is influent with respect to ground water if it contributes water to the zone of satu-

ration. A stream or stretch of a stream is effluent with respect to ground water if it receives water from the zone of saturation (7: 56). The gradient of Bear Creek is greater than that of Horse Creek, and influent stretches along Bear Creek, other things being equal, are shorter and less numerous. The principal influent stretch of Horse Creek is near Lagrange, and it is here that the optimum conditions for irrigation from ground water in the area prevail.

Ground-water recharge from irrigation occurs in two ways: (1) through seepage from canals, and (2) from percolation after use. Mr. C. C. Donahue, water commissioner, states that in the sandy soil of Goshen Hole as much as 65% of irrigation water is lost in transit through percolation and evaporation. No figures on fluctuations of the water table in irrigated areas of this district are available, but there is evidence that the water table is nearer the surface now than before irrigation. In some areas irrigated by the Fort Laramie canal, which diverts water from the North Platte River, this rise in water table is as much as 25 feet.

IRRIGATION SUPPLIES

Successful well irrigation depends largely upon the factors: (1) the water table must be at a relatively shallow depth; (2) there must be a sufficient quantity of water available for irrigation; and (3) the area to be irrigated must be near the well.

At the present time, Mr. Hugh Stemler of Lagrange is the only farmer in the area using well irrigation. His pump, powered by a small farm tractor, lifts water simultaneously from two adjacent wells (nos. 251, 252). Well no. 251 is 33 feet deep and four feet in diameter; well no. 252 is 68 feet deep and two feet in diameter. The pump has a capacity of 1,200 gallons per minute, or approximately 2.7 second feet.

The Quaternary alluvium and gravel here is about 30 feet thick. Both wells obtain the larger part of the water from the alluvium, and the remainder from the underlying White River group. The water level normally fluctuates from six feet below the surface in May to nine feet below the surface in October.

During the 1939 growing season the pumping plant irrigated 160 acres, on which the following crops were grown: beets, 17 acres; beans, 107 acres; barley, 22 acres; corn, 10 acres; and millet, 4 acres. The plant was in operation 96 days. Operating cost, exclusive of depreciation, was \$186.56.

In drilling wells for irrigation, residents are cautioned against (1) drilling wells in localities other than those covered by Quaternary alluvium (Pls. II and III), and (2) drilling deep wells from which water cannot be raised to the surface economically, even if it is encountered in quantities ample for irrigation.

MUNICIPAL SUPPLIES

Hawk Springs and Lagrange are the only villages in the area. Neither has a municipal water-supply system. There are about 12 wells in Hawk Springs, all of which produce from the Lance for-

mation. The wells are about 75 feet deep, and none of them yield water in large quantities; the volume, however, is ample for domestic use. There are about 50 wells in Lagrange. These are shallower than those at Hawk Springs and produce from the White River group and the Quaternary alluvium at a depth of about 35 feet. Water level in the wells is approximately 25 feet below the surface. The supply is more abundant here than at Hawk Springs.

REFERENCES

1. Adams, George I., Geology and water resources of the Patrick and Goshen Hole quadrangles in eastern Wyoming and western Nebraska: U. S. Geol. Survey Water-Supply and Irrigation Paper 70 (1902), 50 pages.
2. Darton, N. H., Blackwelder, Eliot, and Siebenthal, C. E., Laramie-Sherman folio, Wyoming, in Geologic Atlas of the United States: U. S. Geol. Survey Folio 173 (1909), 17 pages.
3. Knight, S. H., and Morgan, Arthur T., Report on the underground water possibilities of the Egbert-Pine Bluffs region: Geol. Survey of Wyoming (1936). Unpublished manuscript in the State Geologist's Office, Laramie.
4., Report on the underground water resources of Crow Creek Valley, Laramie County, Wyoming: Geol. Survey of Wyoming (1937). Unpublished manuscript in the State Geologist's Office, Laramie.
5. Loeffler, Richard J., The stratigraphy of a portion of the east side of the Laramie Range, Laramie County, Wyoming: Geol. Survey of Wyoming (1939), 45 pages. Unpublished manuscript in the State Geologist's Office, Laramie.
6. Meinzer, Oscar E., Ground water for irrigation in Lodgepole Valley, Wyoming and Nebraska: U. S. Geol. Survey Water-Supply Paper 425-B (1917), pp. 37-69.
7., Outline of ground-water hydrology, with definitions: U. S. Geol. Survey Water-Supply Paper 494 (1923), 71 pages.
8. Schlaikjer, Erich M., Contributions to the stratigraphy and paleontology of the Goshen Hole area, Wyoming: Harvard Coll. Mus. Comp. Zoology Bull., vol. LXXVI, no. 2 (1935), pp. 31-68.
9. Theis, Charles V., Amount of ground-water recharge in the Southern High Plains: Am. Geophys. Union Trans., 1937, Part II, pp. 564-568.
10. Surface water supply of the United States: U. S. Geol. Survey Water-Supply Papers 686, 701, 716, 731, 746, 761, 786, 806, and 826.

WELL DATA

DATA FOR WELLS SHOWN ON PLATE II

(Numbers 1-131 inclusive)

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
1	C-35-17-70	J. C. Monday	Wind	6550	17	Dug	9	Qal.	Valley	Gen.	Dry in 1934
2	SW-25-17-70	Harold C. Jones	Hand	6505	62	6"	17	Kp	do	do	Water is very hard
3	NW-25-17-70	John Innes	do	6460	16	Dug	Qal	do	do
4	SE-30-17-69	Chas. F. Smith	Wind	6425	25	do	15	Twr	do	do
5	SE-28-17-69	W. S. Dereemer	do	6450	117	6"	57	do	do	do
6	E $\frac{1}{2}$ -27-69	Nick Contos	Hand	6370	30	Dug	22	Qal	do	do
7	W $\frac{1}{2}$ -13-17-70	John Whitaker	Wind	6560	35	6"	Kp	Hillside	Stock	Not in use
8	NW-8-17-69	Nimmo Livestock Co.	do	7050	110	do	Twr	Draw	do	Never pumps dry
9	SW-4-17-69	do	do	7060	120	do	do	do	do	Not in use
10	C-3-17-69	do	do	6875	110	do	do	do	Stock
11	E $\frac{1}{2}$ -21-18-69	C. W. Hirsig II	Wind	6510	126	do	108	do	do	do
12	NW-17-18-69	do	do	?	do	60	do	do	do
13	6-18-69	do	do	6820	200	do	do	Plains	do	Pumps dry
14	6-18-69	R. J. Welty	do	6840	287	do	do	do	do
15	NE-10-18-69	do	do	6510	440	do	12	do	Valley	Gen.	Tried for flowing well
16	E $\frac{1}{2}$ -21-17-68	Dugald Whitaker	do	6220	30	do	Qal	do	do
17	NE-13-17-68	Nimmo Livestock Co.	do	6120	117	do	60	Twr	Hillside	Stock	Never pumps dry
18	SE-21-18-68	do	do	52	do	do	Valley	do
19	NE-12-19-68	Carl Long	do	6035	143	do	do	Plains	Gen.	Pumps dry
20	S $\frac{1}{2}$ -6-19-67	Claud Long	do	6035	180	do	do	do	do
21	SW-5-19-67	do	do	6015	175	do	do	do	do
22	E $\frac{1}{2}$ -4-19-67	Chas. Cline	do	5860	220	do	do	do	do
23	N $\frac{1}{2}$ -10-19-67	H. W. Gard	do	5820	150	do	do	do	Stock	Very good water
24	SE-35-20-67	Harry Ewart	do	5710	159	do	145	do	Draw	Gen.	Gravel bed produces water
25	SW-36-20-67	G. H. Tate	do	5700	116	do	do	do	do
26	SW-1-19-67	Margaret Hanson	do	5760	201	do	125	do	Plains	do	First water at 93 feet
27	NE-11-19-67	Henry Hanson	do	5760	160	do	130	do	do	do
28	SW-12-19-67	H. W. Gard	do	5740	154	do	135	do	do	do	Can be pumped dry
29	SE-11-19-67	do	do	5820	162	do	do	do	Stock	do
30	E $\frac{1}{2}$ -15-19-67	Frank A. Blas	do	5875	180	do	do	do	Gen.	Pumps dry

DATA FOR WELLS SHOWN ON PLATE II—Continued

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
31	C-14-19-67	W. F. Young	Hand	5730	22	Dug	20	Twr	Plains	Gen.	Pumps dry
32	C-14-19-67	do	Wind	5730	65	6"	do	do	do	Gravel bed produces water
33	E $\frac{1}{2}$ -24-19-67	do	do	5700	38	do	36	do	do	Stock	do
34	E $\frac{1}{2}$ -25-19-67	Percy Laycock	Motor	5700	12	do	Qal	Valley	Gen.	Never pumps dry
35	NW-5-18-67	F. L. DuVall	Wind	5880	75	do	Twr	Plains	do	Can be pumped dry
36	NE-4-18-67	H. A. DuVall	do	5950	40	do	do	do	do	Never pumps dry
37	SE-5-18-67	Ralph Rogers	do	5860	65	do	do	do	do	do
38	10-18-67	Alton Bell	do	5940	100	do	80	do	do	do	do
39	SE-16-18-67	Mrs. Eargle	do	6000	185	do	do	do	do	do
40	NE-19-18-67	Dave Stalker	do	6000	180	do	do	do	do	Has been pumped dry
41	NE-27-18-67	Alex Nimmo	Hand	5940	65	do	do	Valley	do	do
42	NW-29-18-67	Nimmo Livestock Co.	Wind	6025	250	do	150	do	Plains	Stock	do
43	C-1-17-67	J. M. Carey	do	5950	157	do	97	do	Hillside	Gen.	do
44	N $\frac{1}{2}$ -10-17-67	do	do	5950	30	Dug	27	Qal	Valley	do	Gravel bed produces water
45	SW-28-17-67	O. S. Shelton	do	6430	235	6"	205	Twr	Plains	do	do
46	S $\frac{1}{2}$ -26-17-67	Keith Holmes	do	6300	175	do	150	do	do	Stock	Gravel bed produces water
47	S $\frac{1}{2}$ -32-17-67	O. S. Shelton	do	6475	240	do	233	do	do	Gen.	do
48	SE-2-16-67	Keith Holmes	do	6320	160	do	140	do	do	do	do
49	NE-12-16-67	C. S. Vanderpool	do	6280	200	do	163	do	do	do	do
50	SW-6-16-66	Floyd Holmes	do	6270	200	do	180	do	do	do	do
51	N $\frac{1}{2}$ -8-16-66	C. S. Vanderpool	do	6210	178	do	157	do	do	Stock	do
52	SW-2-16-66	John Bevan	do	6150	250	do	do	do	Gen.	do
53	NW-4-16-66	Barney Faber	do	6190	200	do	184	do	do	do	"Sand pumped" every 5 years
54	S $\frac{1}{2}$ -32-17-66	A. Peters	do	6180	156	do	144	do	do	Stock	do
55	C-34-17-66	do	do	6140	215	do	200	do	do	Gen.	do
56	C-14-17-66	Geo. Persae	do	5825	50	do	37	do	do	do	do
57	SE-4-17-66	J. M. Carey	Motor	5740	60	do	30	do	Hillside	Stock	do
58	NW-4-17-66	Mrs. W. P. Owens	Wind	5860	208	do	180	do	Plains	Gen.	do

DATA FOR WELLS SHOWN ON PLATE II—Continued

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
59	NW—5-17-66	Harry Creathbaum	do	5920	156	do	100	do	do	do	
60	SW—32-18-66	Mary Kirkbride	Wind	5930	135	6"	35	Tw	Plains	Gen.	
61	SW—33-18-66	Mrs. Mable Beaver	do	5860	130	do	100	do	do	Gen.	
62	SE—34-18-66	do	do	5780	150	do	100	do	do	do	
63	NW—34-18-66	do	do	5840	130	do	100	do	do	do	
64	NW—32-18-66	Mary Kirkbride	do	5975	224	do	175	do	do	Stock	
65	C—29-18-66	do	do	5920	200	do	165	do	do	do	
66	C—24-18-66	Oscar Pence	do	5600	90	do	do	do	do	do	
67	NW—22-18-66	do	do	do	90	do	do	do	Valley	do	
68	NE—17-18-66	Mary Kirkbride	do	do	60	do	10	do	Draw	do	
69	C—8-18-66	R. Trotter	do	5590	64	do	do	do	Hillside	do	
70	NW—7-18-66	Nimmo Livestock Co.	do	do	120	do	do	do	Draw	do	Never pumps dry
71	NW—11-18-66	Oscar Pence	do	5560	60	do	28	Qal	Valley	Gen.	
72	C—3-18-66	R. Trotter	Hand	5550	12	do	do	Tw	Plains	do	
73	W $\frac{1}{2}$ —5-18-66	Henry Hanson	Wind	6000	130	do	100	do	do	Stock	
74	NE—23-19-66	L. G. Harding	do	5540	120	do	do	do	do	Gen.	
75	C—12-19-66	do	do	5475	140	do	do	do	do	do	
76	N $\frac{1}{2}$ —11-19-66	Earl Marsh	do	5300	70	do	40	do	Valley	do	
77	NW—3-19-66	Peter Schultz	do	5575	120	do	70	do	Plains	Stock	
78	NW—4-19-66	do	do	5615	161	do	90	do	do	Gen.	
79	SW—32-20-66	Max Pensold	do	5600	135	do	100	do	Draw	Stock	First water at 100 feet
80	SE—31-20-66	do	do	5640	165	do	100	do	Hillside	do	
81	S $\frac{1}{2}$ —31-20-66	do	do	5640	173	do	100	do	Plains	Gen.	
82	NW—10-20-65	Chas. Enix	do	5370	160	do	do	do	do	do	
83	N $\frac{1}{2}$ —11-20-65	Bill Wright	do	5410	119	do	112	do	do	do	
84	N $\frac{1}{2}$ —12-20-65	P. J. Peterson	do	5380	139	do	102	do	do	do	First water at 110 feet
85	S $\frac{1}{2}$ —12-20-65	A. M. Marwick	do	5390	124	do	115	do	do	do	
86	E $\frac{1}{2}$ —36-20-65	Jay Bowie	do	5340	130	do	do	do	do	Stock	
87	NW—3-19-65	W. J. Pence	do	5350	145	do	120	do	do	Gen.	
88	NE—9-19-65	do	do	5390	140	do	120	do	do	do	Contact spring
89	N $\frac{1}{2}$ —17-19-65	Ira P. Trotter	Springs	5300	do	do	do	do	Valley	do	
90	SE—13-19-65	Ella Griffin	do	5150	do	do	do	do	do	do	

DATA FOR WELLS SHOWN ON PLATE II—Continued

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
91	N 1/2-10-18-65	H. H. Young	Wind	5575	200	6"	...	Twr	Plains	Gen.	...
92	C-11-18-65	Albert Werner	Wind	5400	200	6"	160	do	do	Stock	...
93	E 1/2-7-18-65	Sam Davis	do	5775	257	do	222	do	do	Gen.	...
94	SE-7-18-65	E. W. Harding	do	5640	165	do	...	do	Valley	do	...
95	S 1/2-9-18-65	Telate G. Werner	do	5460	180	do	100	do	Draw	do	...
96	NW-17-18-65	E. W. Harding	do	5600	150	do	...	do	Valley	do	...
97	C-15-18-65	Albert Werner	do	5460	110	do	80	do	Plains	do	...
98	S 1/2-20-18-65	do	do	5505	140	do	110	do	do	Stock	...
99	S 1/2-35-18-65	Alex Kirkbride	do	5550	100	do	75	do	Hillside	Gen.	...
100	S 1/2-35-18-65	do	do	5445	50	do	25	do	Valley	Stock	...
101	11-17-65	Warren Livestock Co.	do	5940	200	do	...	do	Plains	do	...
102	E 1/2-27-17-65	do	do	5940	222	do	172	do	do	do	...
103	SE-12-16-65	J. C. Berry	do	5910	245	do	232	do	do	Gen.	...
104	SW-4-16-64	D. M. Stevens	do	5750	168	do	145	do	do	do	...
105	35-17-64	James Wilkinson	do	5725	250	do	...	do	do	Stock	...
106	29-17-64	do	do	5825	250	do	...	do	do	do	...
107	19-17-64	do	do	5750	150	do	...	do	do	do	...
108	22-17-64	do	do	5650	260	do	...	do	do	do	...
109	23-17-64	do	do	5700	250	do	...	do	do	do	...
110	12-17-64	do	do	5500	140	do	...	do	do	do	...
111	36-18-64	do	do	5450	100	do	...	do	do	do	...
112	N 1/2-20-18-64	Dan Kirkbride	do	5400	138	do	108	do	do	Gen.	...
113	18-18-64	James Wilkinson	do	...	150	do	...	do	do	Stock	...
114	C-12-18-64	Wallace Powers	do	5300	175	do	160	do	do	Gen.	...
115	NW-27-19-64	Geo. Weaver	do	...	180	do	150	do	do	Stock	...
116	SW-24-19-64	Oscar Yoder	do	5340	180	do	...	do	do	do	...
117	SW-12-19-64	do	do	5010	90	do	30	do	Valley	Gen.	...
118	SE-12-19-64	William Brown	Motor	4900	90	do	10	do	do	do	...
119	SW-8-19-64	W. J. Bunn	Springs	5100	...	11"	...	do	do	do	...
120	NW-4-19-64	Frank Horton	Wind	5340	165	6"	...	do	Plains	Stock	...
121	NE-4-19-64	do	do	5340	200	do	...	do	do	do	...

DATA FOR WELLS SHOWN ON PLATE II—Continued

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
122	C-31-20-64	do	do	5360	170	do	do	do	do
123	NW-32-20-64	do	do	5350	150	do	do	do	do
124	E $\frac{1}{2}$ -33-20-64	Frank Horton	do	5340	200	do	170	do	do	Gen.
125	S $\frac{1}{2}$ -21-20-64	C. D. Fulton	do	5310	212	do	200	do	do	do
126	SW-26-20-64	do	do	5180	226	do	do	do	Stock
127	SE-23-20-64	Arthur Mathoit	Hand	5080	80	do	do	Draw	Gen.
128	S $\frac{1}{2}$ -13-20-64	R. D. Bunn	Wind	4990	20	Dug	12	Qal	Valley	do
129	N $\frac{1}{2}$ -7-20-64	A. J. Ealing	do	5350	190	6"	130	Twr	Plains	Stock
130	S $\frac{1}{2}$ -6-20-64	E. F. Samuelson	do	5340	150	do	90	do	do	Gen.
131	NE-1-20-64	R. E. Swanson	do	5250	240	do	do	do	do	First water at 140 feet

DATA FOR WELLS SHOWN ON PLATE III

(Numbers 132-288 inclusive)

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
132	NE-5-20-63	R. E. Swanson	Wind	5180	180	6"	Twr	Plains	Stock
133	SE-6-20-63	do	do	5220	180	do	do	do	do
134	NW-28-20-63	Geo. Hendrickson	do	4960	135	do	do	do	do
135	SW-27-20-63	do	do	4970	159	do	125	do	do	Gen.
136	SE-28-20-63	do	do	4970	135	do	125	do	do	Stock
137	N $\frac{1}{2}$ -32-20-63	Geo. Parsons	do	4995	145	do	110	do	do	Gen.
138	N $\frac{1}{2}$ -1-19-63	Albert Kessler	do	4845	75	do	do	Valley	do
139	NE-3-19-63	Julia W. Yates	do	4900	167	do	87	do	do	do
140	NE-3-19-63	do	do	4885	15	Dug	10	Qal	Hillside	do
141	N $\frac{1}{2}$ -7-19-63	Arthur Schiack	do	4930	75	6"	40	Twr	Plains	do
142	NE-23-19-63	Pete Haberkorn	do	5100	204	do	200	do	do	do	Small flow into well
143	SW-23-19-63	do	do	5060	170	do	do	Draw	Stock	Never pumps dry

DATA FOR WELLS SHOWN ON PLATE III—Continued

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
144	N $\frac{1}{2}$ —26-19-63	B. A. Hubba	Wind	5140	200	6"	Twr	Plains	Gen.	Pump dry
145	C—25-19-63	F. W. Booth	do	5220	150	do	80	do	do	do	Never pumps dry
146	S $\frac{1}{2}$ —33-19-63	Walter Hunnel	do	5220	108	do	do	do	do	do
147	E $\frac{1}{2}$ —1-18-63	Union Pacific	do	5125	40	do	35	do	Hillside	do	Can be pumped dry
148	E $\frac{1}{2}$ —4-18-63	Harold Hunnel	do	5230	214	do	180	do	Plains	do	do
149	SW—4-18-63	W. J. Hixenbaugh	do	5245	140	do	122	do	do	do	do
150	SE—5-18-63	Alvin Krakow	do	5250	150	do	120	do	do	do	do
151	E $\frac{1}{2}$ —10-18-63	John Donahue	do	5225	175	do	145	do	do	do	Pumps dry
152	C—12-18-63	Clint C. Donahue	5150	400	do	do	do	do	Dry; not in use
153	C—12-18-63	do	5150	200	do	do	do	do	do
154	S $\frac{1}{2}$ —10-18-63	W. F. Hansen	Wind	5165	80	do	68	do	do	Gen.	Can be pumped dry
155	N $\frac{1}{2}$ —14-18-63	Clint Donahue	do	5110	60	do	45	do	do	Stock	do
156	N $\frac{1}{2}$ —14-18-63	do	do	5100	100	do	80	do	do	do	First water at 60 feet
157	C—15-18-63	W. F. Hansen	do	5170	75	do	63	do	do	do	Can be pumped dry
158	NE—20-18-63	John Donahue	do	5330	175	do	145	do	do	do	do
159	NE—22-18-63	Clint C. Donahue	do	5125	65	do	60	do	do	do	Does not pump dry
160	NE—23-18-63	do	do	5080	50	do	30	do	do	do	do
161	N $\frac{1}{2}$ —30-18-63	James Wilkinson	do	5300	100	do	do	do	do	do
162	SW—25-18-63	Clint C. Donahue	Spring	5060	do	do	Valley	Gen.	3/4 sec. ft. flow
163	E $\frac{1}{2}$ —33-18-63	James Wilkinson	Wind	5400	30	6"	Qal	do	Stock	do
164	4 17 63	do	do	5400	75	do	45	Twr	Plains	do	do
165	4 17-63	do	Hand	5300	30	do	20	do	Draw	Gen.	do
166	9 17-63	do	Wind	5500	140	do	do	do	Stock	do
167	18 17 63	do	do	5500	150	do	do	Plains	Gen.	do
168	SW—10-18-62	Alex Shield	do	4920	75	do	do	Valley	do	do
169	SW—9-18-62	Dan Donahue	do	4925	70	do	do	do	do	do
170	W $\frac{1}{2}$ —9-18-62	Meriden Store	do	4970	80	do	45	do	do	do	do
171	NW—11-18-62	Maple Grove Co.	do	4960	75	do	do	do	do	do
172	NE—11-18-62	do	do	4920	75	do	do	do	do	do
173	W $\frac{1}{2}$ —8-18-62	Claud Booth	do	5060	247	do	120	do	do	do	do
174	E $\frac{1}{2}$ —7-18-62	do	do	5105	210	do	47	do	Draw	Stock	Fills from seep at 47 feet
175	E $\frac{1}{2}$ —6-18-62	John W. Brown	do	4980	25	Dug	22	Qal	Valley	Gen.	do

DATA FOR WELLS SHOWN ON PLATE III—Continued

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
176	NE—30-19-62	F. W. Booth.	do	5100	150	6"	Twr	Plains	do
177	SW—18-19-62	Noah Gregory.	do	5100	256	do	90	do	do	do
178	N $\frac{1}{2}$ —18-19-62	E. L. Gregory.	do	4975	80	do	60	do	Draw	do
179	S $\frac{1}{2}$ —7-19-62	Herman Stellflug.	do	5045	300	do	do	do	do
180	S $\frac{1}{2}$ —1-19-62	Leon Lovercheck.	do	4700	280	do	do	Valley	do
181	W $\frac{1}{2}$ —6-19-61	Beat Mt. Store.	Hand	4550	84	do	do	do	do
182	E $\frac{1}{2}$ —2-19-62	O. M. Lovercheck.	Wind	4700	162	do	150	do	Hillside	do
183	N $\frac{1}{2}$ —5-19-62	M. L. Graves.	Motor	4840	69	do	do	Valley	do
184	W $\frac{1}{2}$ —33-20-62	John G. Bunn.	Wind	4760	75	do	do	do	do
185	E $\frac{1}{2}$ —33-20-62	Chas. B. Kessler.	do	4725	35	Dug	do	do	do
186	C—33-20-62	do	Hand	4725	35	6"	do	do	do
187	S $\frac{1}{2}$ —28-20-62	Joe Matje.	Wind	4820	100	do	do	do	do
188	SW—13-20-62	F. M. Smith.	do	4600	60	do	do	Plains	do
189	SE—13-20-62	do	do	4520	28	do	do	do	Stock
190	E $\frac{1}{2}$ —12-20-62	J. A. Test.	do	4560	127	do	100	K1	do	Gen.	Can be pumped dry
191	NW—12-20-62	F. A. Rockhold.	do	4515	198	do	45	do	do	do	Poor water
192	SW—1-20-62	C. M. Rockhold.	do	4490	91	do	84	do	do	do	Good water
193	NW—1-20-62	Elvin Shimm.	do	4440	40	do	do	do	do
194	NE—2-20-62	Joe Fisher.	do	4460	16	do	14	do	do	do
195	C—35-21-62	C. W. Culley.	do	4450	19	do	11	do	do	do
196	SE—26-21-62	Chas. Hinnig.	do	4425	60	do	52	do	do	do
197*	SW—13-21-62	J. H. Swanson.	Hand	4390	72	do	50	do	do	do	Can be pumped dry
198	NE—14-21-62	E. W. Schwab.	do	4360	50	do	35	do	do	do
199	SE—11-21-62	Harry Marlatt.	Wind	4400	32	do	20	do	do	Stock	First water at 11 feet
200	SE—11-21-62	do	Hand	4400	132	do	124	do	do	Gen.	Softer water than No. 199
201	W $\frac{1}{2}$ —12-21-62	G. Korb.	Motor	4360	23	do	16	do	do	do	Does not pump dry
202	NE—11-21-62	Guy Marlatt.	Wind	4380	55	do	do	do	do
203	SW—1-21-62	Earl Woodard	do	4325	14	do	do	do	do	Not in use

* This well is typical of the 12 wells in Hawk Springs, Wyoming. All the wells produce from small seeps. Some wells are slightly shallower and some are slightly deeper than no. 197, but in none of the wells is water very abundant.

DATA FOR WELLS SHOWN ON PLATE III—Continued

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
204	NW—1-21-62	Bob Carlisle	Wind	4315	60	6"	...	Kl	Plains	Gen.	Pumps dry
205	NW—36-22-62	W. A. James	do	4280	150	do	15	do	do	do	First water at 10 feet
206	NE—35-22-62	John Carrelts	Motor	4290	18	do	12	do	do	do	Near creek
207	SW—26-22-62	T. W. Bracken	Wind	4295	208	do	...	do	do	do	...
208	NW—25-22-62	Guy Capek	Motor	4265	135	do	...	do	do	do	...
209	SW—23-22-62	Alex Woods	Wind	4260	255	do	...	do	do	do	...
210	SW—23-22-62	do	Hand	4260	90	do	30	Twr (?)	do	do	...
211	SW—23-22-62	do	...	4260	99	do	30	do	do	do	Not in use
212	NE—11-22-62	Bryan Marlatt	Wind	4245	200	do	...	Kl	do	Gen.	...
213	NE—2-22-62	Kenneth Yorges	Hand	4245	15	do	...	Qal	do	do	Pumps dry
214	N $\frac{1}{2}$ —1-22-62	Henry Wunder	Wind	4240	60	do	40	Twr	do	do	Good water
215	NE—31-23-61	Harry Miller	do	4240	142	do	30	Kl	do	do	Water mineralized
216	NW—32-23-61	Mrs. L. H. Hughes	Motor	4260	82	do	...	do	do	do	do
217	W $\frac{1}{2}$ —34-23-61	James Coxhill	Hand	4240	100	do	30	do	do	do	First water at 50 feet
218	NW—5-22-61	F. M. Hackleman	do	4220	23	do	...	do	do	do	...
219	NW—5-22-61	do	do	4220	30	do	...	do	do	Stock	...
220	NE—6-22-61	A. J. Soller	do	4225	16	do	12	Twr	do	Gen.	...
221	NW—6-22-61	A. E. Springer	do	4225	16	do	...	do	do	do	...
222	NW—6-22-61	do	Wind	4225	16	do	...	do	do	do	...
223	SW—5-22-61	Joe Roushar	do	4220	35	do	...	Kl (?)	do	do	Does not pump dry
224	SE—9-22-61	A. Jones	Hand	4240	26	do	16	Qal	do	do	Flowing; abandoned oil well
225	SE—9-22-61	do	...	4260	5000	2"	...	?	do	Stock	...
226	SW—18-22-61	Edgar Hughes	Hand	4280	18	6"	8	Kl	do	Gen.	...
227	SW—18-22-61	do	Wind	4280	18	do	8	do	do	Stock	...
228	SW—15-22-61	E. W. Talley	do	4280	182	do	100	do	do	Gen.	...
229	NE—23-22-61	T. H. Englehart	Hand	4200	52	do	40	do	do	do	...
230	NW—23-22-61	Hugh J. Bunnell	Wind	4260	152	do	29	do	do	do	...
231	SE—23-22-61	M. Woolam	Hand	4225	23	do	16	Twr	Valley	do	...
232	W $\frac{1}{2}$ —19-22-61	Edgar Hughes	Wind	4300	18	do	8	Kl	Plains	do	...
233	S $\frac{1}{2}$ —19-22-61	John Fricknecht	do	4300	97	do	...	do	do	do	Can be pumped dry
234	N $\frac{1}{2}$ —30-22-61	Ralph W. Hickman	do	4320	125	do	30	do	do	do	First water at 90 feet

DATA FOR WELLS SHOWN ON PLATE III—Continued

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
235	NE-32-22-61	Geo. C. Hacker	do	4340	72	do	30	do	do	do	
236	W $\frac{1}{2}$ -5-21-61	C. A. Harshburger	do	4395	98	do	65	do	do	do	
237	SW-30-21-61	Lincoln Land Co.	Hand	4380	70	do		do	Valley	do	
238	C-32-21-61	G. A. Davis	Wind	4445	100	do		do	do	Stock	
239	C-32-21-61	do	Motor	4470	170	do		do	Plains	Gen.	
240	SW-32-21-61	do	do	4420	18	Dug	8	Qal	Valley	Not in use	
241	SW-31-21-61	do	Wind	4500	70	6'		KI	Plains	Gen.	
242	SE-31-21-61	Lincoln Land Co.	Hand	4500	16	Dug	8	Qal	Valley	Gen.	
243	SE-17-20-61	Lincoln Land Co.	Hand	4445	12	6'	8	Qal	Valley	Gen.	
244	SE-30-20-61	W. A. Gross	Wind	4510	16	do	7	do	do	Stock	
245	NE-32-20-61	do	Hand	4510	20	do	8	do	do	Gen.	
246	NE-32-20-61	do	Wind	4510	20	do	8	do	do	do	
247	SW-33-20-61	Oscar Sherard	do	4520	101	10"	8	Twr	do	do	Tested 1.06 sec. ft.
248	NE-5-19-61	Hugh Stemler	do	4525	52	6'	20	do	do	do	
249	NW-4-19-61	do	Hand	4530	15	do	7	Qal	do	Stock	
250	SE-5-19-61	do	Wind	4545	68	do	20	Twr	do	Gen.	
251a	SW-4-19-61	do	Motor	4570	33	4'	9	Qal	do	Irr.	Back graveled
252a	SW-4-19-61	do	do	4570	68	2'	9	do	do	do	Casing perforated
253b	W $\frac{1}{2}$ -2-19-61	C. F. Moody	Hand	4575	35	Dug	23	Twr	do	Gen.	
254	NE-8-19-61	F. E. Jones	do	4580	70	6'	13	do	do	Not in use	
255	NE-7-19-61	do	Wind	4620	300	do	50	Twr (?)	do	Gen.	
256	C-9-19-61	C. W. Hershey	do	4620	60	do	10	Twr	do	do	
257	N $\frac{1}{2}$ -18-19-61	F. E. Jones	do	4750	56	do	25	do	do	Stock	
258	W $\frac{1}{2}$ -1-18-61	Jelmer Johnson	Springs	4680				do	do	Fill a small reservoir	
259	SE-19-18-60	John E. Latenser	Wind	4920	160	6"	90	do	Plains	Stock	
260	NW-19-18-60	do	do	4830	100	do	90	do	do	Gen.	
261	NW-19-18-60	do	do	4830	205	do	100	do	do	do	Small seep; pumps dry
262	NE-18-18-60	J. H. Draper	do	4820	40	do	36	do	do	do	

a These two wells have a combined output of 1200 gallons per minute. The alluvium and gravel is here about 30 feet in thickness.

b This well is typical of the 50-odd wells in Lagrange, Wyoming.

DATA FOR WELLS SHOWN ON PLATE III—Continued

No.	LOCATION	OWNER	Method of lift	Elev.	Depth	Diam.	Water level	Aquifer	Topog.	Use	REMARKS
263	NE—6-18-60	W. J. Clark	Wind	4800	135	6" ^{am}	100	Twr	Plains	Gen.	
264	SW—31-19-60	Fred McClure	do	4790	139	do	79	do	do	do	Has been pumped dry
265	N $\frac{1}{2}$ —31-19-60	J. P. Curran	do	4760	150	do		do	do	do	Not in use
266	N $\frac{1}{2}$ —32-19-60	Fred McClure	do	4760	190	do		do	do	do	
267	SW—6-19-60	E. L. Chamberlain	do	4700	307	do	170	do	do	Gen.	First water at 300 feet
268	SE—6-19-60	Frank Sanders	do	4690	300	do		do	do	do	
269	SW—6-19-60	E. L. Chamberlain	do	4700	297	do	170	do	do	Stock	
270	NE—6-19-60	Mrs. Sadie Jane	do	4620	300	do		do	do	Gen.	First water at 175 feet
271	SE—19-22-60	J. A. Madden	Hand	4200	87	do	35	Twr (?)	do	do	
272	NE—19-22-60	Alva Stephenson	Wind	4190	167	do	80	do	do	do	First water at 20 feet
273	SW—17-22-60	do	do	4175	115	do	20	do	do	do	
274	SW—18-22-60	Reclamation Serv.	Hand	4130	20	18"	6	Qal	Valley	do	
275	E $\frac{1}{2}$ —16-22-60	Orval Notman	Wind	4160	160	6"		Twr (?)	Plains	do	
276	W $\frac{1}{2}$ —15-22-60	D. G. Greenwald	do	4160	210	do	50	K1	do	do	
277	NW—15-22-60	M. F. Phillips	do	4110	160	do	60	do	do	do	
278	N $\frac{1}{2}$ —15-22-60	Henry Lehman	do	4100	65	do	20	Twr	do	do	Never pumps dry
279	NW—17-22-60	Don Rutten	do	4180	180	do		K1	do	do	Pumps dry
280	SW—10-22-60	A. J. Atkinson	do	4120	120	do		K1 (?)	do	do	First water at 60 feet
281	SE—9-22-60	Ferd Zimmerer	do	4120	166	do	40	do	do	do	
282	SW—8-22-60	Jess Jones	Hand	4110	30	1 $\frac{1}{2}$ "		K1	Valley	do	
283	NW—10-22-60	J. C. Rincker	Wind	4115	160	6"		do	Plains	do	
284	NE—7-22-60	C. D. Sill	do	4120	80	do		do	do	do	
285	SW—5-22-60	J. Yocum	do	4160	150	do		do	do	do	
286	SE—5-22-60	Joe Long	Motor	4140	100	do		do	do	do	
287	SE—3-22-60	C. H. Packer	Wind	4090	200	do		do	do	do	Good water
288	NE—5-22-60	Joe Long	Hand	4130	112	do		Twr (?)	do	do	Has 72 feet of pipe

PIONEER PRINTING CO.
CHEYENNE, WYOMING