

Coal Resources

Between 38 and 65 million years ago, densely vegetated swamps frequently covered what is now Sheridan County. Beneath the waters of these swamps, dying trees, reeds and grasses accumulated as thick layers of dead plant material. As thousands of years passed, sediment was intermittently deposited atop these accumulations, converting them into peat beds. Pressure resulting from the weight of these added sediments compressed the peat, and heat associated with the increasing depth of burial transformed the peat into subbituminous coal. As a result of uplift and subsequent erosion, these coals now crop out over most of the eastern two-thirds of the county. The western third of Sheridan County, however, contains no coal-bearing rocks, partly because of nondeposition and partly because the same erosion that exposed coals in the east completely removed the coal-bearing rocks in the west.

While coal beds are flat-lying or dip very gently westward on the east side of Sheridan County, coals dip eastward in the west central part of the county. Many coals are mapped for more than thirty miles along their outcrops. The lack of detailed mapping, not the lack of coal, accounts for the apparent absence of coal outcrops in some portions of eastern Sheridan County.

The Tertiary Fort Union and Wasatch Formations contain all of the thick and persistent coals that crop out in Sheridan County. In fact, these two formations are two of the most prolific coal-bearing formations in Wyoming. In Sheridan County, the Wasatch Formation contains six to ten persistent subbituminous beds or coal zones. The Uln No. 1 and the Uln No. 2 beds are up to 52 and 30 feet thick, respectively, but most minable coal beds in the Wasatch Formation are between five and fifteen feet thick.

The upper half of the older Fort Union Formation contains up to twelve persistent subbituminous coals, which are a few feet to 45 feet thick. Coals occur in the lower half of the Fort Union Formation as well, but these coals are thinner and much less persistent than those in the upper part of the formation.

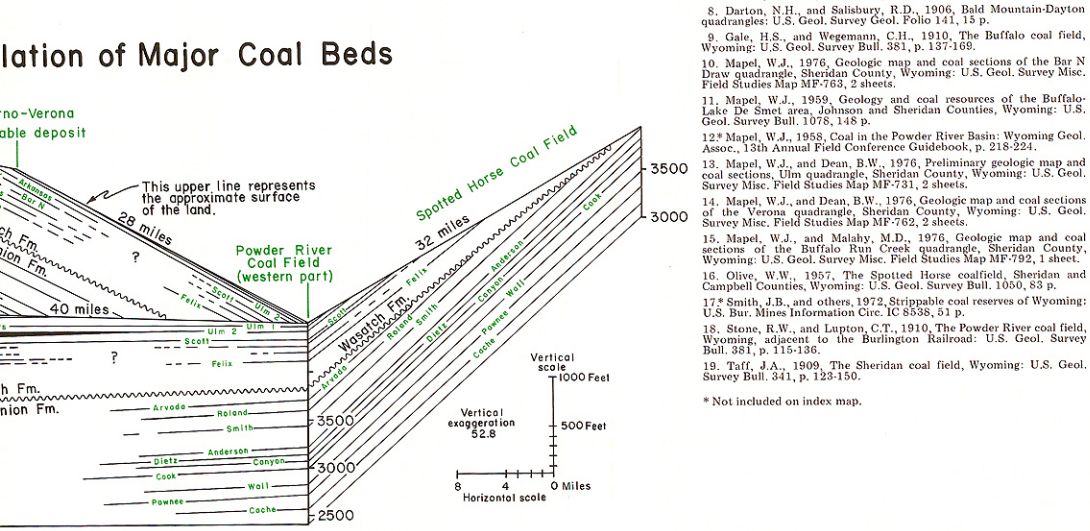
Although the Cretaceous Lance and Mesaverde Formations exhibit some minable low rank coals on the southern and eastern edges of the Powder River Basin, there is no evidence that any but very thin, discontinuous coals occur in these formations in Sheridan County.

Many outcropping coals have been burned through natural causes over large areas of the county. In particular, the thicker coal beds are extensively burned along their outcrops. The heat from these fires has baked and fused overlying rock into distinguishable red "clinker" beds, characteristic of this and many other coal-bearing areas of Wyoming. Despite this burning, more than 26 billion tons of coal still underlie Sheridan County at depths up to 3,000 feet. Of an estimated 7.1 billion tons of coal that occur under less than 1,000 feet of overburden, 5.2 billion tons are believed minable by underground methods.¹ Another 1.9 billion tons that lie under less than 200 feet of rock are potentially recoverable by strip mining.²

Twenty-one percent of Wyoming's remaining identified coal resources and seven percent of its known stripable reserves are in Sheridan County.

¹ Hamilton, P.A., White, D.H., Jr., and Matson, T.K., 1975, The reserve base of U.S. coals by sulfur content, Part 2, The western states: U.S. Bur. Mines IC 8693, p. 267.

² Culbertson and Mapel, 1976, and Smith and others, 1972: see Index and References.

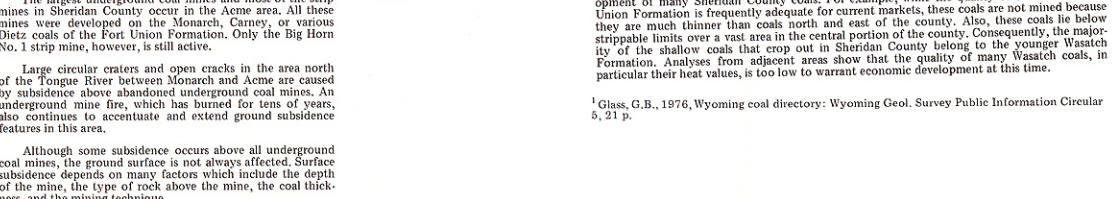
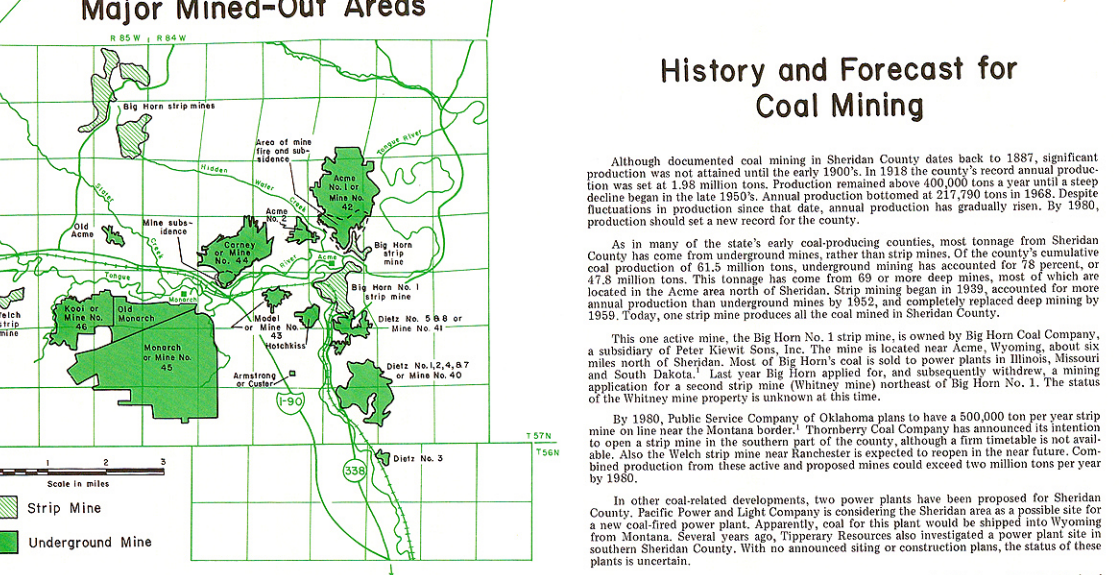
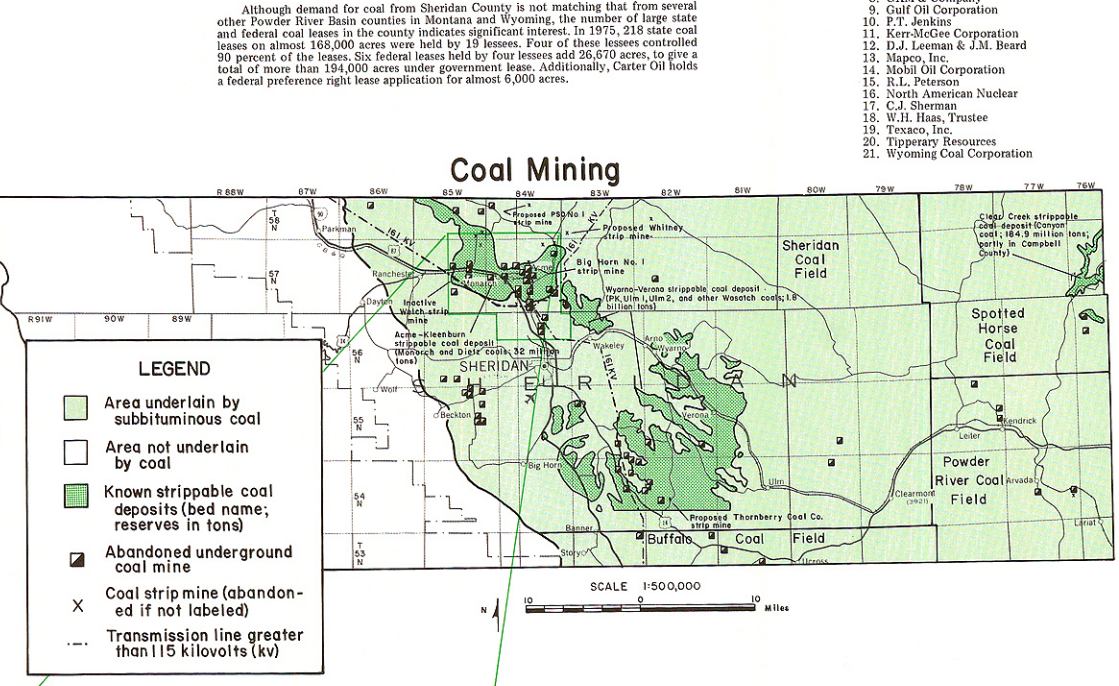
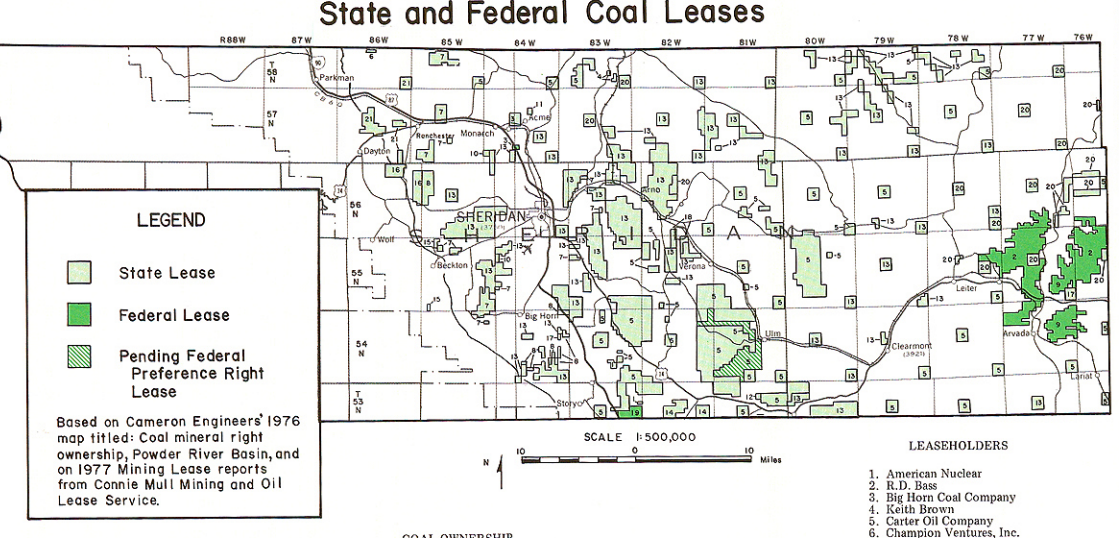


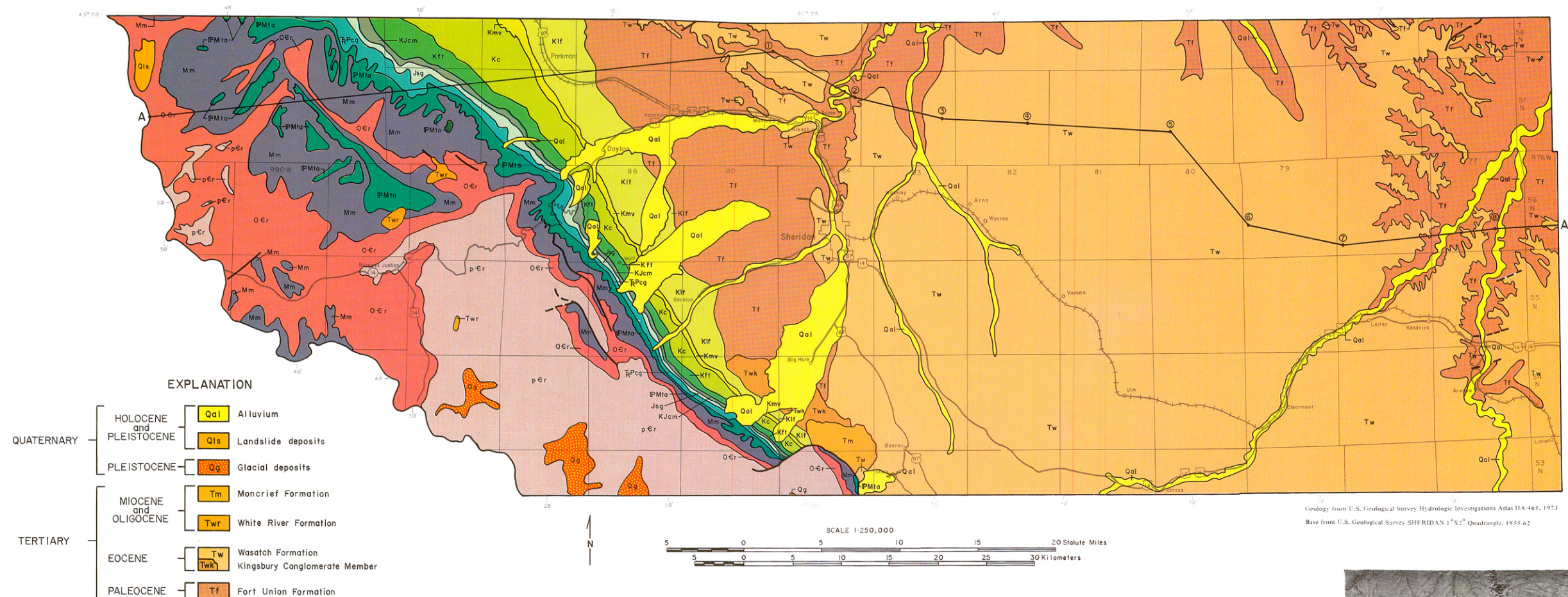
Coal Analyses

FORT UNION FORMATION

	CARNEY		DIETZ NO. 1		DIETZ NO. 2		DIETZ NO. 3		MASTERS ¹		MONARCH		ROLAND ²		SMITH	
	RANGE	AVER.	RANGE	AVER.	RANGE	AVER.	RANGE	AVER.	RANGE	AVER.	RANGE	AVER.	RANGE	AVER.	RANGE	AVER.
AS RECEIVED	46.3 (13) ⁴	46	13	1	14.3 (4) ⁴	14	4	11.3 (2) ⁴	11	2	6.3 (2) ⁴	6	2	76.3 (18) ⁴	76	18
NUMBER OF ANALYSES	46	13	1	14	4	11	2	6	2	76	18	1	1	76	18	1
MOISTURE (%)	21.2-25.8	24.1	23.6	24.7	21.3-23.6	22.4	22.4	19.1-25.7	23.8	21.8	22.7-25.8	24.0	23.3	19.8-25.0	23.0	22.8
VOLATILE MATTER (%)	29.5-35.8	32.0	32.3	37.6	31.3-34.1	33.2	32.8	30.0-34.8	31.6	33.0	32.4-34.8	33.9	34.5	31.2-38.3	33.9	33.4
FIXED CARBON (%)	34.9-43.7	40.3	40.3	33.0	35.5-40.2	37.8	38.1	33.9-41.7	39.5	40.7	36.6-38.2	37.5	37.4	33.1-42.1	38.8	39.2
ASH (%)	2.6-8.6	3.7	3.8	4.7	5.0-8.5	6.7	6.8	3.5-8.6	5.2	4.6	3.8-8.6	4.7	4.9	2.8-11.2	4.4	4.6
HYDROGEN (%)	5.8-6.5	6.3	6.2	6.4	6.1-6.4	6.3	6.3	6.1-6.3	6.2	6.3	6.1-6.3	6.2	6.2	5.8-6.9	6.4	6.5
CARBON (%)	52.3-55.3	53.8	51.5	50.6-53.6	52.2	52.6	52.6	54.5	52.7-53.7	53.1	53.1	53.1	53.1	52.7-53.7	53.9	51.0
NITROGEN (%)	0.9-1.2	1.1	1.1	1.2-1.3	1.1	1.2	1.1	1.3	1.0-1.2	1.1	1.1	1.1	1.1	1.0-1.3	1.1	1.1
OXYGEN (%)	32.5-36.1	34.7	36.1	33.1	31.7-33.1	32.5	32.5	31.3-34.7	33.0	33.0	34.1-34.8	34.5	34.5	30.4-35.0	33.3	34.6
SULFUR (%)	0.3-0.8	0.4	0.4	0.4	0.3-0.8	0.4	0.4	0.3-0.8	0.5	0.5	0.3-0.8	0.5	0.5	0.2-1.2	0.5	0.6
HEAT VALUE BTU/POUND	8910-9220	9240	9260	8903	9000-9430	9200	9200	8730-9170	9060	9360	9010-9360	9230	9250	8450-9360	9350	9410

¹ Undifferentiated Upper and Lower Masters. ² Roland of Taft. ³ Number of analyses used to determine range for proximate analysis, sulfur, and heat value. ⁴ Number of analyses used to determine range for hydrogen, carbon, nitrogen, and oxygen.





Geology from U.S. Geological Survey Hydrologic Investigations Atlas HA-465, 1973.
Base from U.S. Geological Survey SHERIDAN 1°x2° Quadrangle, 1955-62.

STRUCTURAL GEOLOGY

Sheridan County is dominated by two major structural features, the Powder River Basin and the Bighorn Mountains. The Powder River Basin is a structural trough, or syncline, which occupies most of eastern Wyoming. The basin is asymmetric, with its deepest part (axis) lying just east of, and running parallel to, the Bighorn Mountain front. The Bighorn Mountains, a part of which lie in western Sheridan County, are part of a larger, arcuate mountain complex which includes the Pryor, Bighorn, Bridger, and Owl Creek ranges. Structurally, the Bighorns may be described as an asymmetric anticline. As shown on the cross section below, sedimentary rocks lying above the Precambrian gneisses and granites (basement rocks) dip gently or are nearly flat-lying over the top of the range, but dip steeply northeast over the eastern flank of the range where major faulting has occurred.

Five structural elements may be distinguished in the greater Bighorn Mountain region (Hopkin and Jennings, 1971, p. 39). These are northwest trending folds in adjacent basins, the Pryor-Bighorn-Bridger-Owl Creek uplift, east-west trending lineaments, high angle faults and associated monoclinical folds, and flank thrusts along the margins of the uplifts. These elements are considered to have developed in the order given, although considerable overlap exists. For example, many of the northwest trending folds undoubtedly developed contemporaneously with uplift of the adjacent Bighorn Mountains, approximately 55 million years ago.

Lineaments, which are defined as straight or slightly curved features at the earth's surface (faults, breccia zones, etc.), traverse the Bighorn Mountains and adjacent basins at many localities. The Tongue River and Shell lineaments are of interest in Sheridan County (see adjacent photo). The Shell lineament is the surface expression of a Cenozoic mineralized fracture zone, a part of which may be the northern tear fault of the Piney Creek thrust (NW $\frac{1}{4}$, T.53N., R.84W.) The Tongue River lineament, trending northeast across the Bighorns in alignment with the Tongue River, is characterized by faulting and mineralized fracture zones. Faulting of a segment of the lineament along the north end of the Precambrian core in Sheridan County has resulted in

at least 600 feet of stratigraphic displacement. A northeast extension of the Tongue River lineament has been recognized in the subsurface of the Powder River Basin.

Flank thrusts are locally exposed along the eastern margin of the Bighorns, as shown on the geologic map (e.g., T.53N., R.84W.). These are loosely called thrusts, but seismic evidence suggests that these flank faults steepen at depth and have only a minor component of horizontal displacement at the surface. Foster and others (1968, p. 100) have postulated a near-vertical subsurface fault extending from the Casper Arch in Natrona County north to Sheridan, with a maximum throw (vertical displacement) of 4000 feet at Buffalo. A similar subsurface fault along the west flank of the Bighorns (near Basin) suggests vertical movement of the Bighorn Mountain block(s).

The Sheridan County region has been subjected many times to tectonic activity (mountain building, basin subsidence, etc.), but Cenozoic tectonism is responsible for most of the present large-scale structural features. The distribution and ages of Cenozoic strata along the Bighorn Mountain front are important clues to unraveling the tectonic history of the area. The oldest Cenozoic gravel deposit, the Kingsbury Conglomerate Member of the Wasatch Formation, contains pebbles and cobbles of Paleozoic and Mesozoic rocks. Hopkin and Jennings (1971, p. 43) believe that the Kingsbury represents the first significant and widespread uplift of the range.

Tectonic movement of the Piney Creek thrust (T.53N., R.84W.) during the Miocene is indicated by conglomerates of the Mancosch Formation. The Mancosch, containing cobbles and boulders of Precambrian rocks, is distributed along the front of the Piney Creek thrust. Pliocene uplift of the Bighorns is indicated by Oligocene strata lying unconformably on Paleozoic strata high in the Bighorn Mountains (T.56N., R.88 and 89W.). These Oligocene outcrops have been vertically displaced several thousand feet above equivalent strata in the adjacent Powder River Basin in latest Tertiary time. Possible Quaternary uplift of the Bighorns is indicated by

gravel-capped pediments and terraces adjacent to the east flank of the Bighorns; many of these increase in gradient directly adjacent to the mountain front, suggesting very recent vertical movement. However, the possibility of climatic control on the origin of these terraces must be considered.

For additional stratigraphic and geological information, see the Stratigraphy Plate.

REFERENCES

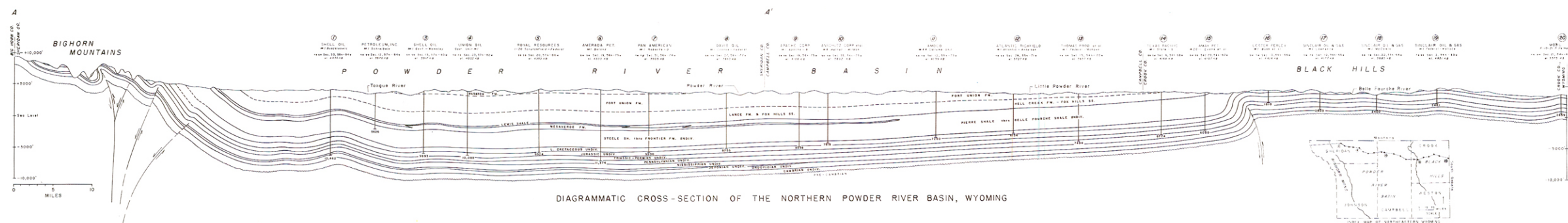
- Darton, N.H., 1906, Geology of the Bighorn Mountains: U.S. Geol. Survey, Prof. Paper 51, 129 p.
- Demarest, M., 1941, Critical structural features of the Bighorn Mountains, Wyoming: Geol. Soc. America Bull., v. 52, p. 161-176.
- Foster, N.H., Goodwin, P.E., and Fisher, R.E., 1968, Seismic evidence for high-angle flank faulting, Bighorn Mountains, Wyoming: Geol. Soc. America, Abs. for Ann. Meeting, p. 100-101.
- Hudson, W.G., Pearl, R.H., and Druse, S.A., 1973, Water resources of the Powder River Basin and adjacent areas, northeastern Wyoming: U.S. Geol. Survey Hydrologic Investigations Atlas HA-465.
- Hopkin, R.H., and Jennings, T.V., 1971, Cenozoic tectonic elements, Bighorn Mountain region, Wyoming-Montana: Wyoming Geol. Assoc. 23rd Ann. Field Conf. Guidebook, p. 39-47.
- Hudson, R.F., 1969, Structural geology of the Piney Creek thrust area, Bighorn Mountains, Wyoming: Geol. Soc. America Bull., v. 80, p. 283-296.
- Osterwald, F.W., 1959, Structure and petrology of the northern Bighorn Mountains, Wyoming: Wyo. Geol. Survey Bull. 48, 47 p.
- Wyoming Geological Association, 1976, Geology and energy resources of the Powder River: 28th Annual Field Conference Guidebook, 328 p.



BIGHORN MOUNTAINS. The aerial photograph shows the Bighorn Mountain uplift, flanked to the west by the Bighorn Basin, and to the east by the Powder River Basin. Dark areas of the photograph represent vegetation. The western part of Sheridan County is outlined by a dashed line. Sedimentary rocks dipping steeply into the Powder River Basin are clearly shown along the east flank of the mountains. On the west side of the mountains, these same strata dip gently into the Bighorn Basin.

Several structural features are visible on the photograph. The Pryor Mountains are the less uplifted northern part of the Bighorn Mountains which extend into Montana. The Tongue River lineament is clearly shown as a northeast trending zone which traverses the range. The Shell lineament traverses the range in the southern part of Sheridan County, and, according to Hopkin and Jennings (1971, p. 42), is partly indicated by the linearity and landslide topography of Shell Canyon. The Piney Creek thrust is shown as a dramatic eastward bridge along the east margin of the Bighorn Mountains. According to Hudson (1969, p. 283), the thrust has moved a sledge of Precambrian and Paleozoic rock 2.5 miles eastward from the main mountain front, resulting in a stratigraphic displacement of 10,000 feet.

Many drainage and other physiographic features show up well on the photograph. The interested reader may correlate features shown on the photograph with those on the geologic map. FROS Data Center photo.



A land-use map depicts the prevailing uses of natural and man-altered land, including both urban and rural uses. It does not reflect the optimum or potential use of a given area.

A map of an area as large as Sheridan County, compiled and checked within the span of one year, cannot be free of error. This map represents a "best estimate" of predominant land use in the summer of 1976.

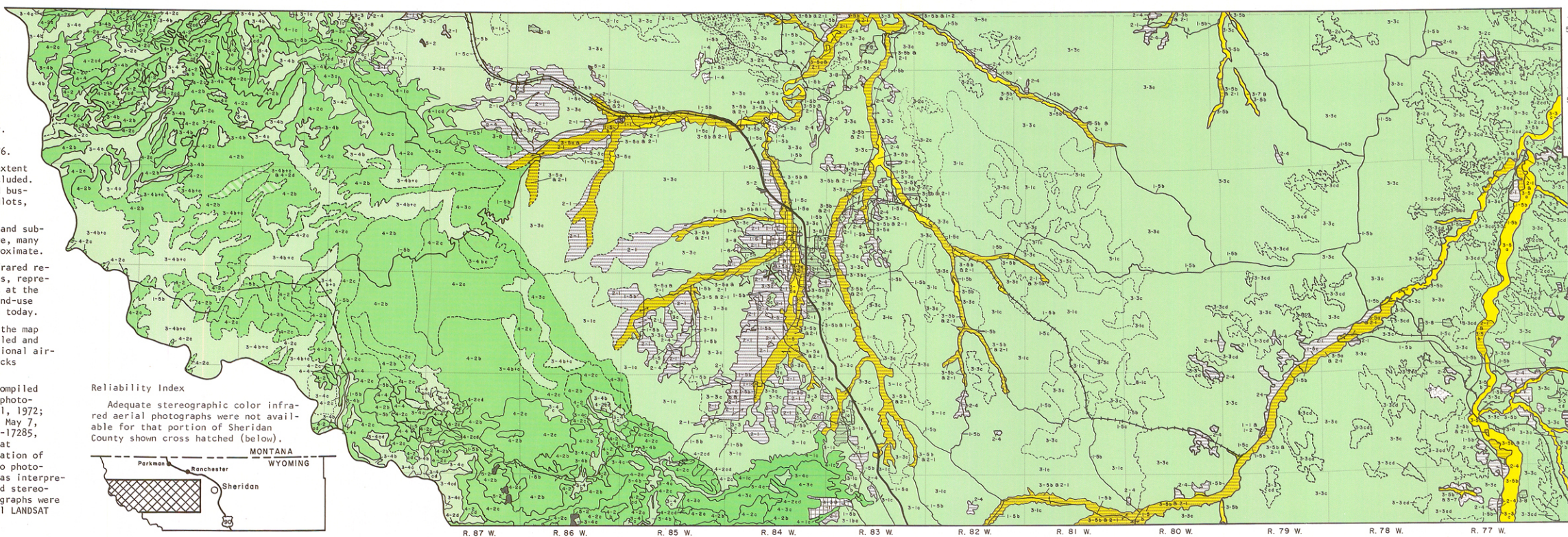
Only those land-use classes of areal extent appropriate to the map scale could be included. Excluded are isolated urban dwellings and businesses, ranch houses and buildings, feedlots, small stockpounds, and the like.

Many land-use units (especially grassland subclasses) grade into one another; therefore, many of the mapped boundary locations are approximate.

Land-use identifications, based on infrared reflectance of vegetation and other features, represent the conditions of the ground surface at the time of overflight. In areas of rapid land-use change, such identifications may not hold today.

Most of the airphotos used to compile the map were obtained in 1973. The map was compiled and revised in 1976-77, on the basis of additional airphotos obtained in May 1975 and field checks carried out during the summer of 1976.

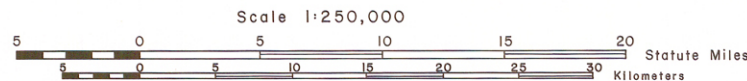
The Sheridan County Land-Use Map was compiled from high-altitude color infrared aerial photographs (NASA flights no. 72-138, August 11, 1972; no. 73-147, August 30, 1973; and no. 310, May 7, 1975) and from LANDSAT band 7, scene 1409-17285, September 5, 1973. The map was compiled at 1:100,000-scale by stereo-photo interpretation of 1:120,000-scale photographs. Where stereo photographs were unavailable, LANDSAT band 7 was interpreted at 1:250,000-scale. Areas not covered stereographically by high-altitude aerial photographs were transferred directly from the more general LANDSAT interpretations.



Land use map compiled by: D.R. Gaylord, Department of Geology, University of Wyoming.

Work supported by: U.S. Geological Survey, Grant #14-08-0001-G-163.

MAP KEY



URBAN AND BUILT UP LAND:

- 1-1 Subclass Residential: High-density, extensive areas of habitation, such as Sheridan and Ranchester. Individual ranch and farm sites not mapped.
- 1-2 Subclass Commercial: High-density, extensive areas of commercial structures, places of business and small industry.
- 1-4 Subclass Extractive: All extractive enterprises such as open pit coal mines and oil and gas wells. Open pit coal mines are designated by solid lines. Oil and gas production areas, rather than being mapped as individual wells, were mapped as "fields".
- 1-5 Subclass Transportation: Major maintained county, state and federal highways; 1-5a four-lane interstate highways; 1-5b all two-lane highways; 1-5c railroad tracks.
- 1-6 Subclass Recreation: Recreation areas including golf courses, athletic fields, picnic and camping areas, and fairgrounds.

AGRICULTURAL LAND :

- 2-1 Subclass Irrigated and Sub-irrigated Crop-land Pastureland: Agricultural land (harvest and/or pastureland) which has been periodically maintained or improved by artificially supplying water or by using naturally high groundwater supplies in flood plains to grow native hay, alfalfa, etc. Dual use of irrigated land and flood plain is common.
- 2-4 Subclass Dryland Crop: Crops dependent on rain and/or snowfall for their water needs. These croplands quite often can be identified by the parallel strips of fallow and cultivated land.

RANGELAND :

- 3-1 Subclass Moist Shrub Grassland: Areas of lush shrubs and deciduous trees which predominate in the low spots and valleys. The suffixes b and c (good and fair, respectively) which are based on near-infrared reflectivity indicate the quality of range.
- 3-2 Subclass Semiarid Shrub Grassland: Most common in areas of drier range where dry shrubs, coniferous trees and bushes are interspersed with range flora. The suffixes b and c (good and fair, respectively) which are based on near-infrared reflectivity indicate the quality of range.
- 3-3 Subclass Grassland: Encompasses vast areas of land use on the semiarid prairie lands. In rating the quality of range an arbitrary classification was employed based on near-infrared reflectivity of the soils: the higher the reflectivity,

(Rangeland - continued):

the lower the range classification. The highest reflectivities were found in areas of active and near-active erosion. The smoother, more rolling plains received higher classifications. Those with the least dissection and sufficient water received the highest rating.

- 3-3b Good rangeland.
3-3bc Good to fair rangeland.
3-3c Fair rangeland.
3-3cd Fair to poor rangeland.
3-3d Poor rangeland.

NOTE: In most cases the boundaries between grassland types can only be approximately located.

- 3-4 Subclass Mountain Meadows: Rangeland at higher elevations bordered on three or more sides by dense coniferous growth. This range is predominantly used for summer pasture of sheep and cattle brought up from the warmer, drier lowland range. Subclasses are based on reflectivity, topography and lushness of growth. The areas with substantial sagebrush growth received a lower rating.
- 3-4b Good mountain meadow, low relief, low erosion with a substantial supply of moisture.
- 3-4c Fair mountain meadow, high relief predominating, higher erosion with numerous exposed barren areas and a lesser supply of moisture. This rangeland may have up to 10% coniferous cover over a total area. The conifers may occur in small dense growths or as separately spaced trees.
- 3-4cd Fair to poor mountain meadow--a mixed classification with small areas of 3-4c and 3-4d intermingled.
- 3-4d Poor mountain meadow, barren rock and talus in many places with sparse lichen growth, grasses and small shrubs. This type generally occurs above the tree line in a zone next to barren rock.

3-5 Subclass Natural Flood Plain: The flood plain is defined as the area periodically inundated under high water conditions. However, the area denoted as flood plain on this map includes more infrequent high water marks also.

- 3-5a Flood plain with greater than
deciduous cover.
- 3-5b Flood plain with less than 25%
deciduous cover.

NOTE: The flood plain may also support a variety of shrubs; but, differences in shrub types are not denoted in the classification.

(Rangeland - continued):

- 3-7 Subclass Naturally Disturbed Vegetation: Vegetated land which has been subjected to changes in ground and/or surface waters. These areas contain numerous barren soils which often are the result of adverse alkalinity or too great an influx of recent sediment. Consequently this range subclass often exists as a mixed land-use type with flood plain and near flood plain areas.
- 3-8 Subclass Previously Cultivated: Land which was formerly under cultivation, showing no signs of recent agricultural activity. It is apparently being allowed to return to a natural state.

FOREST LAND:

- 4-1 Subclass Deciduous: Denotes the densest growths of deciduous trees which primarily occur in and along the flood plains and within urban districts.
- 4-2 Subclass Coniferous: Coniferous forest areas which have been subdivided on the basis of crown density. Without specific ground-truth statistics this cannot be expected to accurately depict the health and/or productivity of the forest.
 - 4-2b Densest conifer cover (greater than 90%) with few open meadow spaces or barren rock exposures.
 - 4-2c Moderately dense conifer cover (greater than 60%). The remainder is open meadow and/or barren rock with limited use as rangeland.
 - 4-2cd Mixed areas of 4-2c and 4-2d.
 - 4-2d Least dense conifer cover (10-30%). This forest type is often found along steep-sided, talus-strewn valleys. These areas may have limited use as rangeland.
- 4-3 Subclass Mixed: Areas of greater than 10% deciduous cover coexisting with conifers.
 - 4-3c Moderately dense mixed forest (greater than 60% conifer density). This subclass has limited use as rangeland.
 - 4-3cd Mixed forest (30-60% combined crown density).
- 4-4 Subclass Clear-cut: All areas cleared of timber in the recent past for lumbering purposes. A dual use at this time is as mountain meadow because of the abundance of range flora at many of the clear-cut sites.

WATER:

- 5-2 Subclass Lakes and Reservoirs: Large natural lakes, man made lakes and reservoirs.

NONVEGETATED LAND:

- 7-3 Subclass Barren Rock: Includes exposed crystalline or sedimentary bedrock, talus fields, or other exposed modern sediments. These barren rocks have less than 5% organic cover including grasses, shrubs or trees. The most extensive areas of barren rock occur above the tree line.

KEY TO OTHER MAP SYMBOLS :

- Boundary of land-use unit
 ✶ ✶ Airstrip
 - - - - - Oil or gas field boundary



SOURCES OF DATA

Anderson, J.R., Hardy, E.E., and Roach, J.T., 1972, A land-use classification for use with remote sensed data: U.S. Geological Survey Circular 671, 16 p.

Marr, R.W., 1975, Special report: Land-use in the Moorcroft and Keyhole Reservoir areas, Crook County, Wyoming: University of Wyoming, Department of Geology, August, 1975, 17 p.

Wyoming Geological Survey, 1976, Johnson County, Wyoming—geologic map atlas and summary of land, water and mineral resources: Wyoming Geological Survey, County Resource Series No. 4, 9 maps.

A clear-cut area (4-4) with abundant wood debris and new seedling growth lies in the foreground. Fair conifer forest (4-2c) is visible at the far clear-cut boundary. Poor conifer cover (4-2d) and fair-to-poor mountain meadow (3-4cd) cap the barren, rocky ridge on the horizon. The picture was taken looking NE from NW $\frac{1}{4}$ sec. 25, T.54N., R.88W.

Fair rangeland (both 3-3c and 3-1c) can be seen in the foreground with deciduous trees (4-1) growing on the low sloping flanks of the mountains. Poor and fair conifer cover (4-2c and 4-2d, respectively) as well as fair mixed deciduous and conifer forests (4-3c) are visible on the steep face of the mountain. The picture was taken looking SW from NW $\frac{1}{4}$ sec. 23, T.54N., R.85W.

A mountain lake (5-2) is set amongst moderately dense (4-2c) and dense (4-2b) conifer forest. The barren rock outcrops (7-3) seen above the lake are not large enough to be included on the map. The photo was taken looking NE from NE $\frac{1}{4}$ sec. 35, T.54N., R. 87W.

Fair rangeland (3-3c) in the foreground gives way to a dryland field (2-4) on the level ground immediately below the butte. Fair to poor rangeland (3-3cd) occupies the upper levels of the butte with fair rangeland (3-3c) covering the flanks. This picture was taken looking N from NW ¼ sec. 4, T.54N., R.76W.

An open pit coal mine (1-4) is surrounded by fair rangeland (3-3c). The Bighorn Mountains are visible on the horizon. The photo was taken looking S from NW $\frac{1}{4}$ sec. 36, T.58N., R.85W.



Irrigated cropland (2-1) and flood plain (3-5b) are seen in foreground and center. Fair to poor rangeland (3-3cd) covers the slopes of the uplands in the background. The picture was taken looking SE from SE $\frac{1}{4}$ sec. 30, T.54N., R.79W.

The landforms map represents a classification of the land surface of Sheridan County. Each area is classified according to its dominant geomorphic characteristics. A landform unit is defined on the basis of regional and local topography, pattern, texture, and the geomorphic processes that control its formation. Boundaries between landform units should represent the lines of sharpest contrast between landform types; however, because some boundaries are not abrupt, their placement is uncertain, as indicated on the map by dashed lines.

The Sheridan County landforms map was compiled from interpretations of Skylab color infrared photographs (S-190A, track 5, pass 10, June 13, 1973, roll 15, frames 229-231) and high-altitude color infrared aerial photographs (NASA flights no. 72-138, August 11, 1972; no. 73-147, August 30, 1973; no. 239, June 20, 1973; no. 248, August 7, 1973; and no. 310, May 7, 1975). The map was first compiled at 1:120,000-scale by stereo photointerpretation of aerial photographs. This interpretation served to define general landform classes and mappable units. The photo interpretation was then field checked, revised as necessary, and recomputed. This recomputation served to fill in areas obscured on the aerial photographs by clouds and to improve the accuracy of map-unit boundaries. Areas not covered stereographically by high-altitude aerial photographs (see reliability index) were interpreted from Skylab photographs.













Map Key

Landforms Modified by Man:




-  Surface modified for urban development.
-  Surface modified by mining activities.

Erosional Landforms (forms created by the destructive action of denuding factors):

Landforms of fluvial origin:



-  D-1 Widely-spaced dendritic drainage with large, flat interstream areas; interstream divides indefinite.
-  D-2 Dendritic drainage with intermediate spacing; interstream areas steeper and narrower than D-1; interstream divides readily defined.
-  D-3 Closely-spaced dendritic drainage with many small, steep interstream areas; interstream divides readily defined.
-  T-1 Widely-spaced trellis drainage with large, flat interstream areas; interstream areas indefinite.
-  T-2 Trellis drainage with intermediate spacing; interstream areas steeper and narrower than T-1; interstream divides readily defined.
-  T-3 Closely-spaced trellis drainage with many small, steep interstream areas; interstream divides readily defined. Smaller drainage enters larger drainage at right angles.
-  2-1 Knobby topography showing high degree of irregular dissection and dominated by small buttes.
-  2-2 Surfaces near base of slopes currently reworked by fluvial processes.
-  3-1 Steep slopes and irregular topography cut by steep-sided stream valleys.
-  3-2 Dip-slopes. Sloping surface conforms closely to the dip of the supporting rock layers.
-  3-3 Broad, sloping mountain valleys; valley floors composed of alluvial and colluvial material.
-  Gr Granitic terrain; irregular, rounded topography with some bare rock summits.

Landforms of fluvio-glacial origin:


-  GI-1S Rounded glacial topography with rounded valleys, ridges, and knolls marginal to steeper glacial topography; some bedrock exposed.
-  GI-1 Glacier-scoured topography with bedrock exposed in U-shaped valleys, on steep valley walls, and on broad divides.
-  GI-2 Steep, glacial topography with bedrock exposed in aretes, horns, cols, and steep-sided glacial valleys.

Depositional Landforms (forms created by the aggradational action of water, wind, ice, or mass movement):


Deposition by water:

-  4-1 Flood plains of major streams and rivers; level bottom-land where alluvium has accumulated.
-  S-1 Smooth, gentle slopes flanking major streams and rivers.







Deposition by ice:

-  6-1 Moraines and glacial outwash material forming hummocky, low-dipping topography that extends outward from regions of glacial erosion; currently being reworked by fluvial processes.

Deposition by mass wasting:

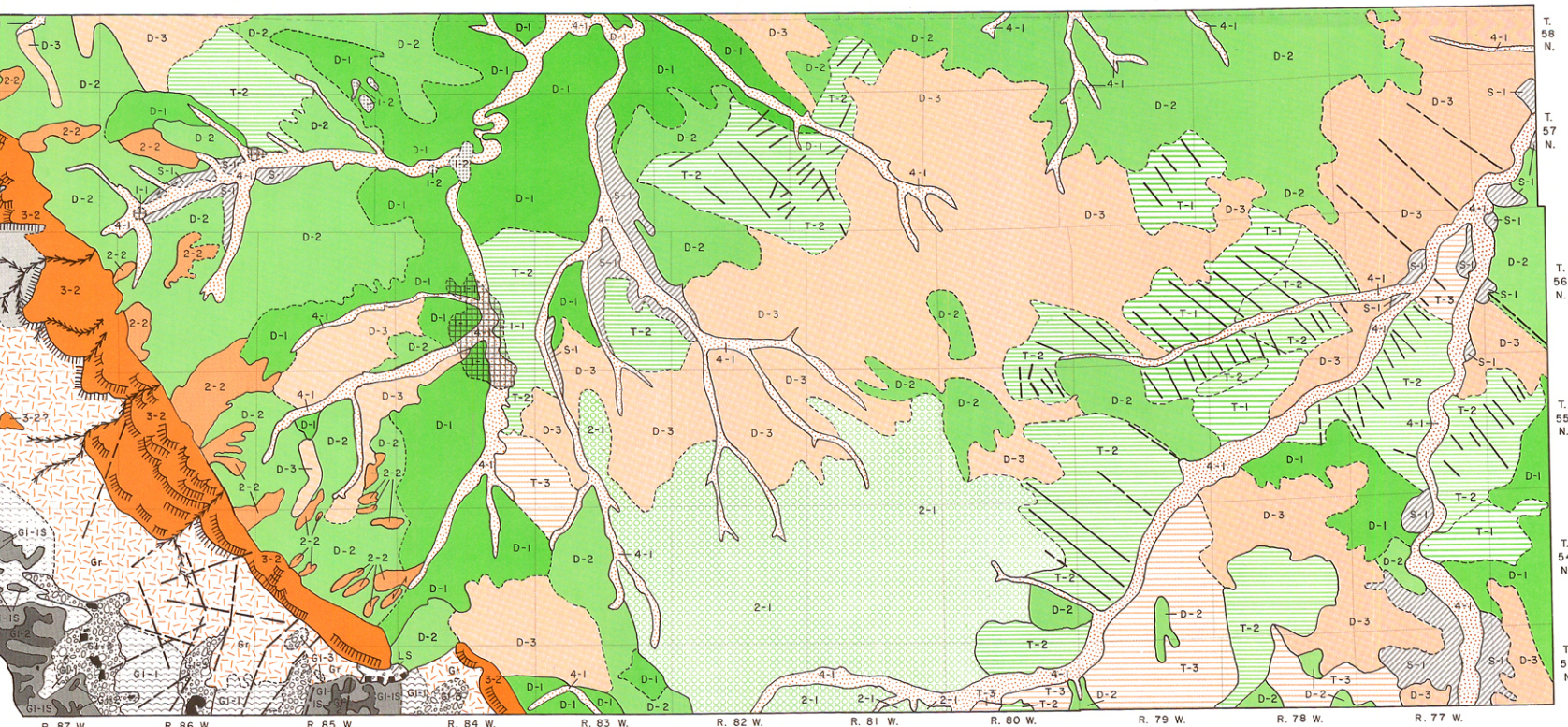
-  LS Large-scale landslides, earth-slumps, and mud-flows resulting in hummocky terrain beneath steep, breakaway scarps.

Key to Map Symbols:

-  Boundary of landform unit; dashed where uncertain
-  Narrow, steep-walled canyon; arrows point downstream
-  Abrupt change in slope; hachures on steeper slope
-  Cliff or near-vertical slope; blocks on cliff side
-  Large-scale linear feature; dashed where less obvious
-  Body of water

Reliability Index

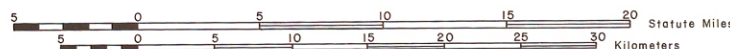
Stereoscopic color infrared aerial photographs were not available for that portion of Sheridan County shown cross hatched below.



Landforms map of Sheridan County compiled by:
R.W. Marrs, D.R. Gaylord, and J.K. King
Department of Geology, University of Wyoming

Work supported by:
The U.S. Geological Survey
Grant 14-07-0001-G-163

Scale 1:250,000



View looking southeast from SE $\frac{1}{4}$ NE $\frac{1}{4}$ sec. 27, T. 58N., R. 88W. A typical flood-plain area flanked by rangeland occupies the central portion of the scene. Outcrops of older sedimentary rocks form eastward-dipping dipslopes in the right background.



View looking northeast from U.S. Highway 14 toward Dayton. Large, flat pediment surfaces slope eastward away from the mountain front (left center).



View of a large slump area, shown looking west-southwest from the southwest corner of sec. 6, T. 54N., R. 84W. The hummocky ground at the base of the escarpment is typically unstable.



View of a broad mountain valley in western Sheridan County, sec. 14, T. 55N., R. 90W. This rolling topography is typical of much of the area along the crest of the northern Bighorn Mountains.



An area of recent surface collapse caused by the burning-out of coal from beneath the surface. This underground coal fire is near Acme.

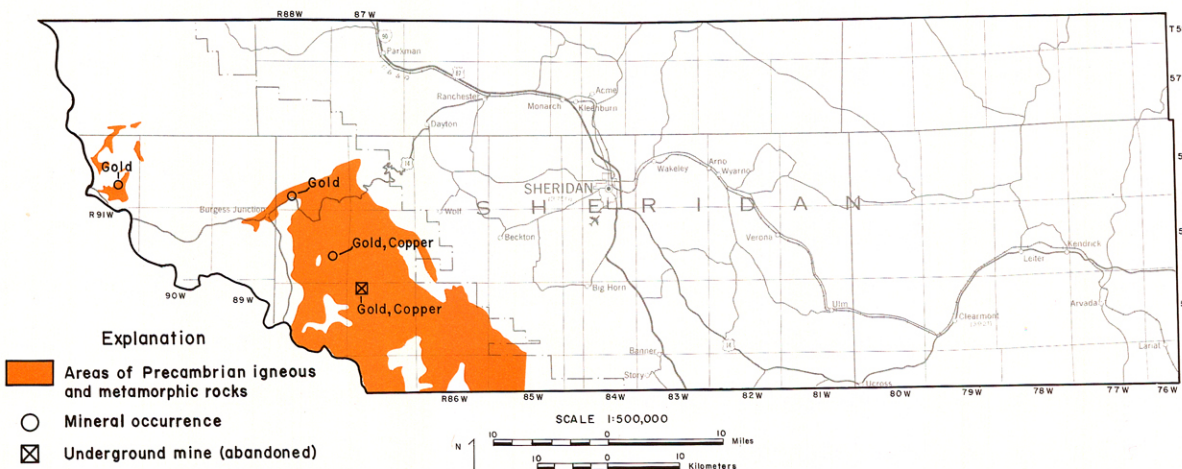


View of Sawmill Lakes, looking northeast from NW $\frac{1}{4}$ sec. 35, T. 54N., R. 87W. This area is an excellent example of glacial terrain with lakes formed on a hollow which was "carved-out" by moving ice and surrounded by moraines left as the glaciers melted.

SOURCES OF DATA

- Cooke, R.U., and Doornkamp, J.C., 1974, Geomorphology in environmental management: Oxford University Press, London, 413 p.
- Curran, H.A., Justus, P.S., Perdew, E.L., and Prothero, M.B., 1965, Atlas of landforms (2nd. ed.): John Wiley and Sons, New York, 140 p.

Minor Minerals



- Explanation**
- Areas of Precambrian igneous and metamorphic rocks
 - Mineral occurrence
 - Underground mine (abandoned)

Minor Minerals

Copper

Copper mineralization occurs at a number of small prospecting pits and abandoned mines on the east flank of the Bighorn Mountains. The most common copper minerals in this region are malachite, which occurs primarily as apple-green staining in fractures in diabasic and granitic rocks, and chalcocite, which occurs as massive, metallic, lead-gray crystals in fractures. Some minor amounts of galena, a metallic, lead-bearing mineral, and chrysocolla, a blue-green hydrous copper silicate, occur in association with some of the copper veins.

Gold

Placer gold deposits have been found near the headwaters of the Little Bighorn River (Beeler, 1907). The gold was reported to occur as fine flat grains with sharp jagged edges, suggesting little abrasion during stream transport.

Gold, in lode deposits, is generally associated with some silver in quartz veins in granitic and diabasic rocks along shear zones and at contacts between dikes and the intruded granitic masses.

The basal sandstone of the Flathead Formation is known to contain gold at several sites in the Bighorn Mountains. However,

attempts to mine the gold have proven unprofitable.

The interested reader is referred to Osterwald and others (1966) for additional information on gold prospects in Sheridan County.

Iron

Iron mineralization is very minor in Sheridan County. The common iron-bearing minerals magnetite, hematite, pyrrhotite, pyrite, and limonite are generally associated with copper mineralization and occur as massive and disseminated minerals in shear and fracture zones.

At the old "Leopard Rock" quarry south of Twin Buttes, stringers of disseminated bronze-colored pyrrhotite and massive pyrite occur in diabasic dikes and on the old mine dumps.

Sheridanite

Sheridanite is a talc-like variety of chlorite. It was named after Sheridan County, where it was first found and identified (Osterwald and others, 1966). Sheridanite has no economic significance.

Tungsten

Scheelite has been reported in granite in prospect pits and

shallow shafts about 25 miles southwest of Sheridan, where microscopic tungstite and wolframite are reported to occur in the quartz-rich rocks.

Vermiculite

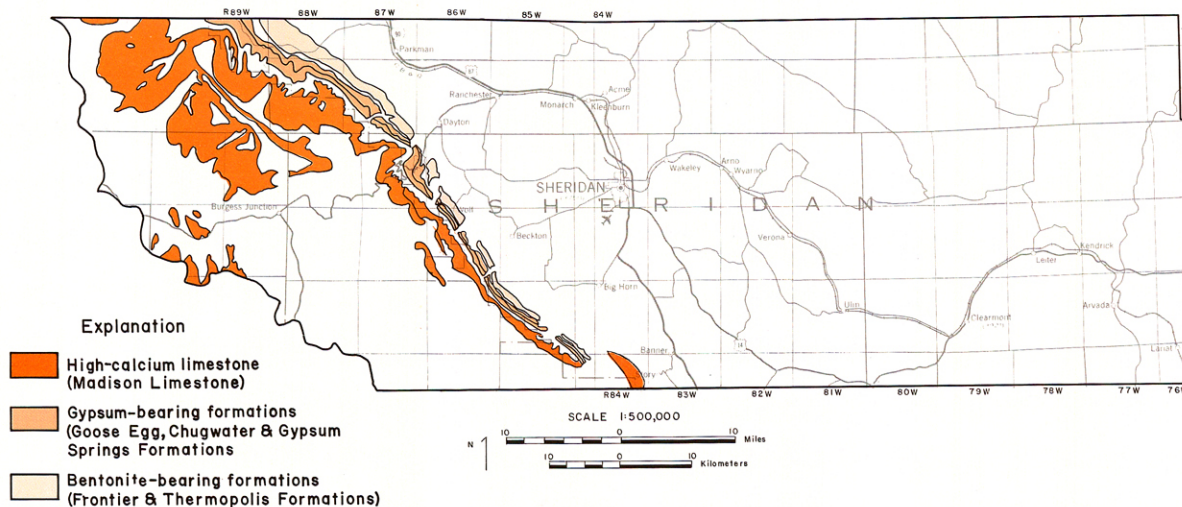
Vermiculite is a soft micaceous magnesium-aluminum-iron silicate which has the unique property of expanding to up to 30 times its original volume when heated or acid treated. Because of this property, vermiculite is commonly used for insulating purposes.

Vermiculite occurs about 30 miles southwest of Sheridan in a number of prospecting pits; however, no production has been realized (Hagner, 1944).

References

- Beeler, H.C., 1907, unpublished report no. 103 for the Geological Survey of Wyoming, 7 pp.
- Hagner, A.F., 1944, Wyoming vermiculite deposits: Geological Survey of Wyoming Bulletin 34, 47 pp.
- Osterwald, F.W., Osterwald, D.B., Long, J.S., Jr., and Wilson, W.H., 1966, Mineral resources of Wyoming: Geological Survey of Wyoming Bulletin 50, 287 pp.

Industrial Minerals



- Explanation**
- High-calcium limestone (Madison Limestone)
 - Gypsum-bearing formations (Goose Egg, Chugwater & Gypsum Springs Formations)
 - Bentonite-bearing formations (Frontier & Thermopolis Formations)

Bentonite

The term, bentonite, was first used by Wilbur C. Knight (1898) for clayey material in the Cretaceous Benton Formation of eastern Wyoming. Since that time, bentonite has been recognized for its unique property of increasing in volume from 15 to 20 times when wetted (Hosterman, 1973, p. 127).

Two kinds of bentonite are recognized, the sodium-rich clay, or swelling variety, which is known as "western" or "Wyoming" bentonite, and the nonswelling calcium-rich bentonite known as "southern" bentonite. (Gillson, 1960).

Bentonite typically forms by devitrification and chemical alteration of glassy igneous rock such as tuff and volcanic ash. Cretaceous formations containing bentonite beds today were deposited from airborne volcanic ash in Wyoming during a time of extensive volcanism in northern Idaho and western Montana. Bentonite is typically light green on fresh fractures and alters to a cream or white color when exposed to the air, and has a greasy soapy feel.

Bentonite beds in Sheridan County occur in the Cretaceous Thermopolis, Mowry and Frontier Formations, which are exposed along the east flank of the Bighorn Mountains. Commercial deposits of bentonite have been identified in northern Sheridan County (T.58N., R.88W.) by Knechtel and Patterson (1956, p. 99). These deposits have been named the Clay Spur, Bed J, and Soap Creek beds. The Clay Spur lies within the Mowry Formation, the Clay Spur bed has been measured at 5 1/2 feet thick, and is well suited for use as foundry and drilling mud clay. Bed J, 3 1/2 to 6 1/2 feet thick, is equally well suited for commercial use. The Soap

Creek bentonite is up to 20 1/2 feet thick and was commercially stripmined in sec. 17 and 18, T.58N., R. 88W., by the Wyoming Bentonite Mining Company in the late 1940's. However, thick overburden has precluded extensive bentonite mining in this area.

In order to have commercial production of bentonite, the following criteria must be met: the bentonite seam must be nearly horizontal to gently dipping (less than 5 degrees), the overburden cover must be thin, the bentonite seam must be of a minable thickness and of good quality, and there must be adequate means of transportation to the market to help minimize costs.

Bentonite is used in foundries for heavy duty casting, and as a binder for the pelletization of powdered iron-ore concentrate in the taconite industry. The petroleum industry uses bentonite in drilling muds and as a catalyst in petroleum refining. It is also used as a carrier for insecticides and fungicides and as a seal and waterproofing material for irrigation canals and reservoirs (Hosterman, 1973).

The processing of bentonite includes the following steps, weathering, drying, grinding, sizing, granulation, and the addition of additives for cation exchange (see Hosterman, 1973).

Gypsum

Gypsum is a naturally occurring sulfate mineral. It is typically deposited in evaporitic sedimentary environments in association with limestone, salt, marl, and clay. Gypsum, $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$, contains 79 percent calcium sulfate and 21 percent water, by weight. The transparent variety of gypsum is called selenite, while the white or lightly colored massive variety is called alabaster. Other varieties of gypsum include satin spar and gypsite.

The Goose Egg, Chugwater and Gypsum Spring Formations contain significant deposits of gypsum. In Sheridan County, these formations form a northwest-trending outcrop band on the east flank of the Bighorn Mountains. Gypsum is continuously exposed at the base of the Chugwater Formation from the Wyoming-Montana border south to Little Goose Creek (Stone, 1920, p. 299).

Gypsum is a common building material. It is used in wallboard and plaster. Crude gypsum is used in portland cement to retard setting, and in agriculture to neutralize alkaline soils and provide sulfur. Calcined gypsum ($\text{CaSO}_4 \cdot 1/2\text{H}_2\text{O}$) is marketed as plaster and used in wallboard (Reed, 1975).

The location of gypsum mines with respect to markets is a critical factor affecting their economic feasibility. Marketing gypsum deposits are well situated with respect to markets in the Pacific Northwest, an area deficient in gypsum deposits. However, the steep dips of gypsum-bearing formations in Sheridan County may preclude extensive development of gypsum mining there.

High-Calcium Limestone

A third industrial mineral of potential economic value in Sheridan County is high-calcium limestone. High-calcium limestone is already in demand in Wyoming for use in refining sugar from sugar beets and in lime scrubber systems for air pollution control of coal-fired power plants. In these power plants, the high-calcium limestone is used to remove sulfur from stack gases.

The Madison Limestone (Mississippian) is the best potential source of high-calcium limestone in Sheridan County. The Madison is variable in thickness, up to 800 feet thick at the Little Tongue River (Sando, 1976, p. 47). The Madison is underlain disconformably by the Jefferson Dolomite (upper Devonian) and overlain disconformably by the Amsden Formation (Late Mississippian and Pennsylvanian), which lies on a karst surface formed on the Madison in Late Mississippian time.

The Madison has been divided into six members (Sando, 1976) on the basis of rock type and faunal content. These are the Cottonwood Canyon Member (basal member consisting of silty and sandy dolomite with subordinate shale and quartz sandstone), the Lower Hurst Member (limestone, dolomitic limestone, and dolomite), a cherty dolomite member (poorly fossiliferous, fine-grained cherty dolomite), a cherty limestone member (medium- to thick-bedded, cherty, crinoidal limestone), and the Bull Ridge Member (cherty, coarse-grained, fossiliferous limestone). Of these members, parts of the Woodhurst, cherty limestone, and Bull Ridge may contain beds of high-calcium limestone.

A mine was proposed for the Madison Limestone west of Story in Sheridan County, approximately one mile north of the Johnson County Line. However, environmental considerations have curtailed development. The limestones at the Story outcrops are very pure, ranging up to 98.7 percent pure calcium carbonate. A potential obstacle to limestone mining in Sheridan County is the relatively steep dip (averaging 30 degrees east) of the Madison Limestone along the flank of the Bighorn Mountains. Steep dips generally require underground mining for economic feasibility.

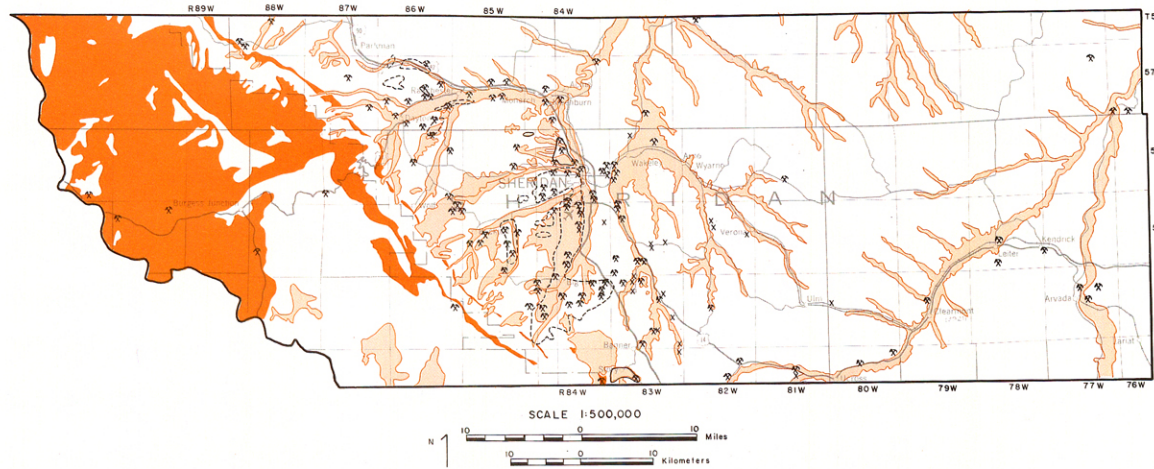
In the event of large-scale construction of new coal-fired electrical generating plants in the Powder River Basin, it is likely that high-calcium limestones will be used for scrubbers. As a rough approximation, about 100,000 tons of limestone would have to be mined each year to supply a 1000 MW generating plant, depending

on variables such as heat value of the coal and purity of the limestone. High-calcium limestone beds in Sheridan County are an important resource in this regard.

References

- Gillson, J.L., 1960, Bentonite, in Industrial minerals and rocks: A.I.M.E. Seeley W. Mudd Series, A.I.M.E., Gillson, J.L., ed., New York.
- Hosterman, J.W., 1973, Clays, in United States Mineral Resources: U.S. Geological Survey Professional Paper 820.
- Knechtel, M.M., and Patterson, S.H., 1956, Bentonite deposits in marine Cretaceous formations, Hardin district, Montana and Wyoming: U.S. Geological Survey Bull. 1023.
- Knight, W.C., 1898, Bentonite: Engineering and Mining Journal, v. 66, p. 491.
- Reed, A.H., 1975, Gypsum: U.S. Bureau of Mines Bull. 667, p. 1-9.
- Sando, W.J., 1976, Madison Limestone (Devonian and Mississippian), east flank of Bighorn Mountains, Wyoming: Wyoming Geological Association 28th Annual Field Conference Guidebook, p. 45-52.
- Stone, E.W., 1920, Gypsum deposits of the United States: U.S. Geological Survey Bull. 697.

Construction Materials



Explanation

- Gravel pits
- Scoria pits
- Excellent gravel area delineated by sampling and testing.
- Good to fair gravel deposits.
- Unconsolidated surficial deposits, includes alluvium, terrace, pediment and glacial deposits.
- Potential sources of crushable ledge rock material.

Compiled from Wyoming Highway Department Construction Materials Survey and listing of known sand and gravel pits.

Sheridan County Construction Materials Summary

SYSTEM	SERIES	FORMATION	THICKNESS	CHARACTER	CONSTRUCTION MATERIAL POTENTIAL
QUATERNARY	Recent	ALLUVIUM-COLLUVIUM	0-60'	Sand, silt, clay, and gravel; sediments are fine grained and may be from the Big Horn River; also includes landslide and lake deposits.	Best the mountain gravel can be found in these deposits. In most cases, however, the water table is at or near the ground surface.
		TERACE DEPOSITS	0-50'	Clay, silt, sand and gravel. Gravel consists of well-sorted pebbles, cobbles and subrounded boulders of limestone, sandstone, quartzite, chert and igneous rocks. In southern part of county may have 0-5' of silt, sand and clay overburden.	Best source of highway construction material in this county. Gravel is often clean, quite hard and well-sorted.
	Recent and Pleistocene	PEDESTAL DEPOSITS	0-100'	Heterogeneous mixture of Paleozoic and igneous rocks, grading related to distance from mountains.	Not a good source of construction material. Contains large boulders, as well as interstitial clay. An extensive screening and crushing program would be required.
		WINDSETT CONGLOMERATE	10-1200'	Lower one-third interbedded greenish siltstone, silty sandstone, thin beds of conglomerate grading upward to boulder beds.	Could be a source of highway construction material but contains some bentonitic material.
TERTIARY	Pliocene	WINDSETT CONGLOMERATE	10-1200'	Light brown bentonitic siltstone interbedded with calcareous sandstone and light gray tuffaceous sandstone.	This formation generally will not provide any construction material due to interbedding of shale causing high plasticity, only the scoria beds are recommended for use in highway construction.
		WHITE RIVER FORMATION	10-300'	Shale, sandstone, bentonite, and lignite layers having an overall drab brownish-gray appearance in upper part. Some white sandstone and conglomerate in lower part.	Could be a source of highway construction material but contains some bentonitic material.
	Oligocene	WASATCH	500-2400'	Interbedded sandstone, siltstone, and shale.	This formation generally will not provide any construction material due to interbedding of shale causing high plasticity, only the scoria beds are recommended for use in highway construction.
		KIMBERLY CONGLOMERATE	10-600'	Randomly conglomerate, some sandstone and claystone, consists mainly of hard Paleozoic rocks.	Could be a source of highway construction material but contains some bentonitic material.
CRETACEOUS	Paleocene	FOOT HILLS	2000-4000'	Interbedded clay, shale, sandy shale, and coal in lower and upper parts; middle part consists of massive sandstone beds.	Sandstone beds would yield nonplastic sand and sandstone.
		LANCE	1500-3000'	Brownish sandstone, gray shale, and carbonaceous shale; some lignite in upper part. Typically sandstone in a cross-bedded channel deposit.	May have some sand and sandstone, but generally contains too much shale and is not recommended for use.
		FOX HILLS SANDSTONE	100-800'	Salt and pepper sandstone; contains gray shale at top and massive, cliff-forming reddish-brown sandstone at bottom.	A satisfactory source of sand and sandstone, but not hard enough for crushable ledge rock.
		HEADWATER	200	Reddish sandstone that weathers buff and blue.	No construction materials located within this formation.
	Cenozoic	PEASBERRY	600-700'	Upper and middle parts consist of carbonaceous shale and silty sandstone. Peasberry sandstone member (lower part) is massive, cross-bedded, and medium grained.	Peasberry sandstone member could supply unlimited quantities of sand and sandstone, but not hard enough to provide crushable ledge rock.
		GOOSE EGGS	3500-3600'	Medium to dark gray marine shale; contains sandstone zone in upper part.	Although parts of this formation are sandy, the formation should be avoided where interbedded shales tend to cause high plasticity.
		FRONTIER SANDSTONE	400-800'	Interbedded sandstone and shale at the top from the Hell Creek Sandstone member, consisting of three sandy, bentonitic beds at the base of the formation.	The two top Hell Creek Sandstone members could provide nonplastic sand and sandstone, but the lower Hell Creek member is not recommended, nor is any part of the lower Frontier Formation recommended for any use due to its bentonite content and high plasticity.
		MOWRY SHALE	200-350'	Dark brownish gray silty shale that weathers to silty gray; contains thin beds of bentonite.	This formation is hard, brittle, and shatters in place; it is not recommended for use in highway construction. The material tends to slip and is only recommended as "gravel surfacing" for non-paved roads.
	JURASSIC	THEMOPOLIS SHALE	150-250'	The upper part consists of gray shale and beds of bentonite. Lower part consists of thin-bedded black shale. The Muddy Sandstone member is a well-indurated drab buff to brown, medium-grained sandstone.	No construction material deposits located within this formation.
		CLARKVILLE	40-100'	Massive white to light-brown medium-grained sandstone containing lenses of small pebbles conglomerate (10-30') with 100' of grayish-brown shale.	The sandstone member could provide nonplastic sand and soft sandstone, but is not hard enough for crushable ledge rock.
		NEVADIAN	300-350'	Interspersed beds of fine to medium-grained sandstone and shale, weathers to variegated colors.	No construction material deposits located within this formation.
		SHENANDO	250-300'	Interbedded sandstone, siltstone, and shale; light-gray and yellow sandstone with 3-12' interbedded limestone layers at top, middle and bottom.	The Shenando sandstone is an excellent crushable ledge rock. It will have dip slopes available with several hundred thousand cubic yards available.
TRIASSIC	TRIASSIC	CHOCOMA	750-1000'	Bright red, fine to medium-grained sandstone, shale, and siltstone containing gypsum beds as much as 2 feet thick.	No construction material deposits located within this formation.
		GOOSE EGGS	350-250'	Soft, reddish-brown, silty clay and siltstone, interbedded gypsum and red siltstone.	No construction material deposits located within this formation.
		NEVADIAN	300-400'	Massive, light-gray, white, yellow, or pinkish-white fine to medium-grained cross-bedded sandstone.	This formation is an unlimited source of non-plastic sand and sandstone, but has only a few thin layers of shale and is not recommended for use as a crushable ledge rock.
		ANDER	250'	Red sandstone and claystone and tan massive dolomite.	The sandstone and clay are not a source of construction material, but the dolomite layers may provide a crushable ledge rock.
	MISSISSIPPIAN	MADISON Limestone	500'	Thin-bedded to massive resistant lime dolomite and limestone unit forms prominent cliffs.	This formation has an unlimited supply of limestone with qualities meeting the most rigid demands of high quality materials specifications.
		SHENANDO	150-300'	Dark gray to red sandstone overlain by resistant pitted gray dolomite. Top layer is lenticular, shaly dolomite.	The dolomite is recommended as an excellent source of crushable ledge rock. It would be hard and would meet high quality specifications.
		GALATHEA Limestone	50'	Grayish-red, shaly, silty limestone, with some fine pebbles conglomerate.	This formation is not recommended because it is often obscured by overburden; overlying units have better quality material in open exposures.
		CHOCOMA	500'	Soft non-resistant medium to coarse-grained green to red sandstone overlain by soft grayish-green shale.	This formation is often covered by overburden and vegetation. It is not recommended for use as a construction material.
	CAMBRIAN	FLATHEAD SANDSTONE	360'	Thin to brown, medium to coarse-grained quartz sandstone with some conglomerate at base.	Could provide nonplastic sand and sandstone, but no hard ledge rock is available.
		CHOCOMA	500'	Massive, light-gray and gray, fine to coarse-grained granitic intrusive rocks of varying composition and metamorphic rocks.	Could provide crushable hard rock as construction material; may also have some of diagenetic granite which might be used without blasting.
		CHOCOMA	500'	Massive, light-gray and gray, fine to coarse-grained granitic intrusive rocks of varying composition and metamorphic rocks.	Could provide crushable hard rock as construction material; may also have some of diagenetic granite which might be used without blasting.
		CHOCOMA	500'	Massive, light-gray and gray, fine to coarse-grained granitic intrusive rocks of varying composition and metamorphic rocks.	Could provide crushable hard rock as construction material; may also have some of diagenetic granite which might be used without blasting.

Compiled from Wyoming Highway Department Construction Materials Survey

Construction Materials

Sheridan County has a number of geologic units suitable for construction materials, particularly in the western part. Most often the location of a deposit with respect to its use determines whether it is economic. Energy related development of coal, uranium, and oil in the Powder River Basin will increase the demand for construction materials in Sheridan County.

The best sources of sand and gravel in the county are terrace deposits, most of which are located near the Bighorn Mountains or along the major stream drainages. Alluvial stream deposits become progressively finer away from the mountains and are generally too fine for construction uses.

Crushable aggregate sources are large and include Precambrian, Ordovician, Mississippian and lower Jurassic rock units. Most of the Upper Cretaceous and Tertiary units are too soft for such use.

Clinker, or baked rock formed from natural underground coal fires, is found in abundance where the Fort Union and Wasatch formations are exposed. Clinker is used for ballast and subsurfacing or surfacing on secondary roads; it is generally brittle and has a high absorption rate.

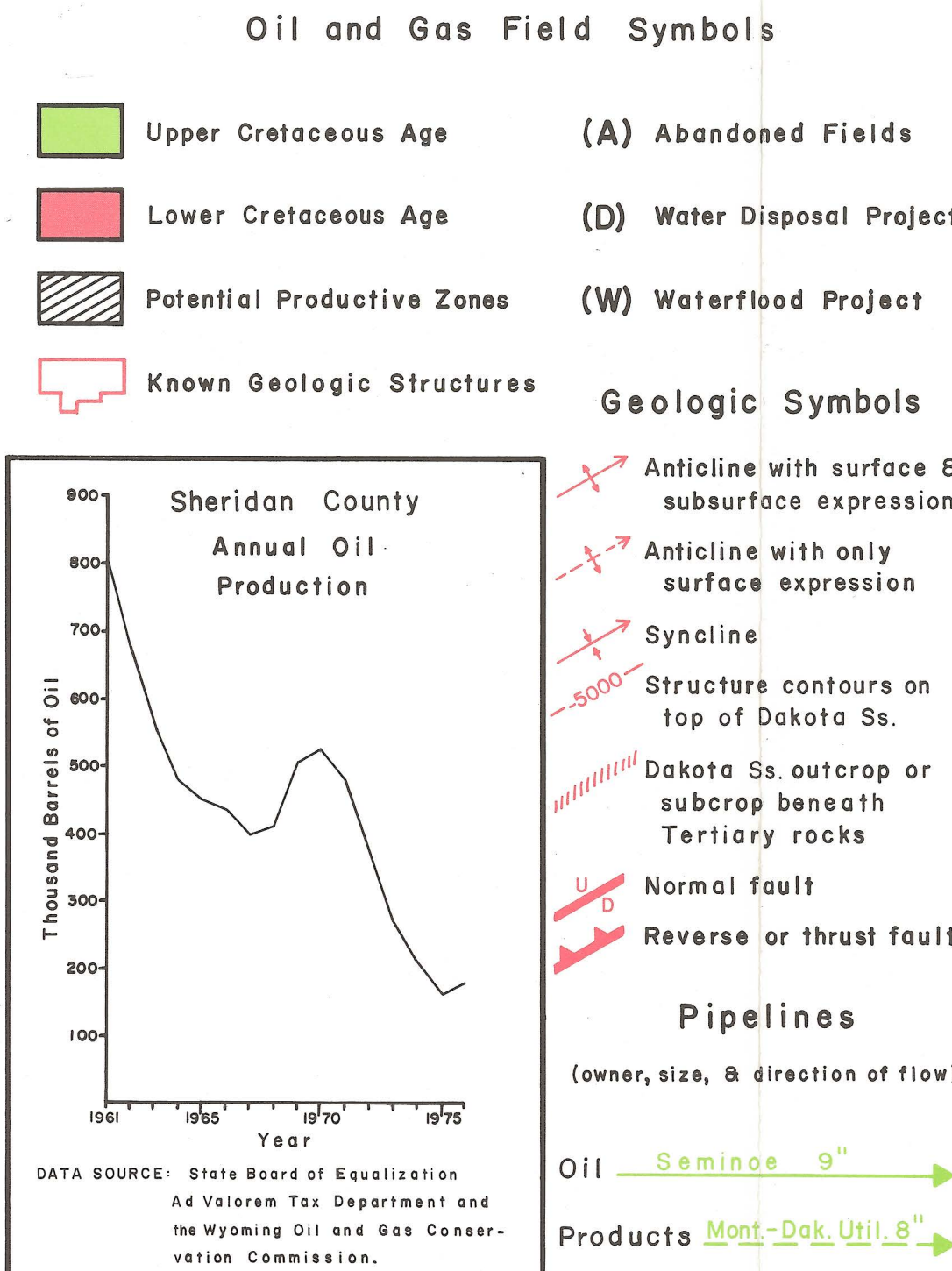
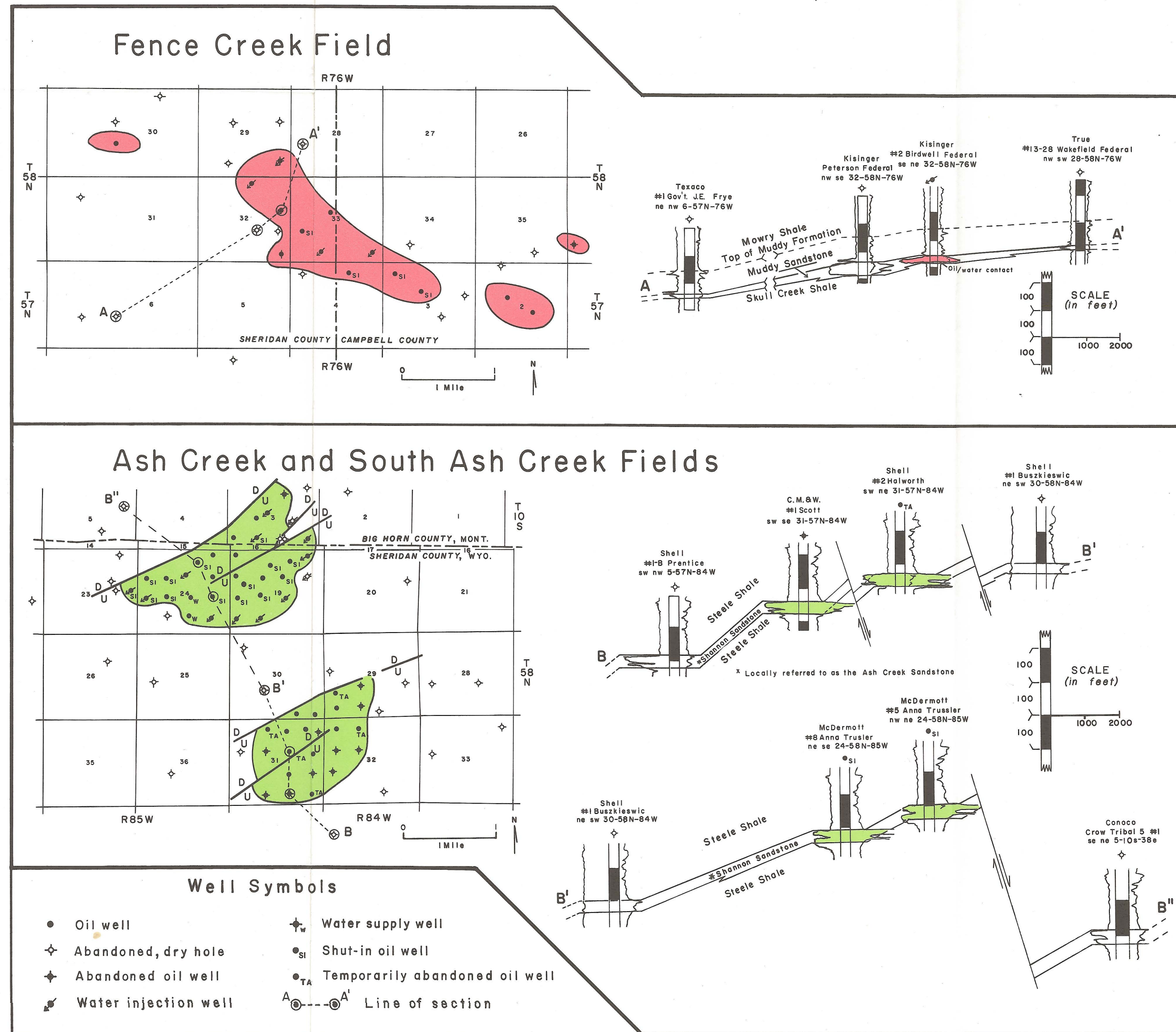
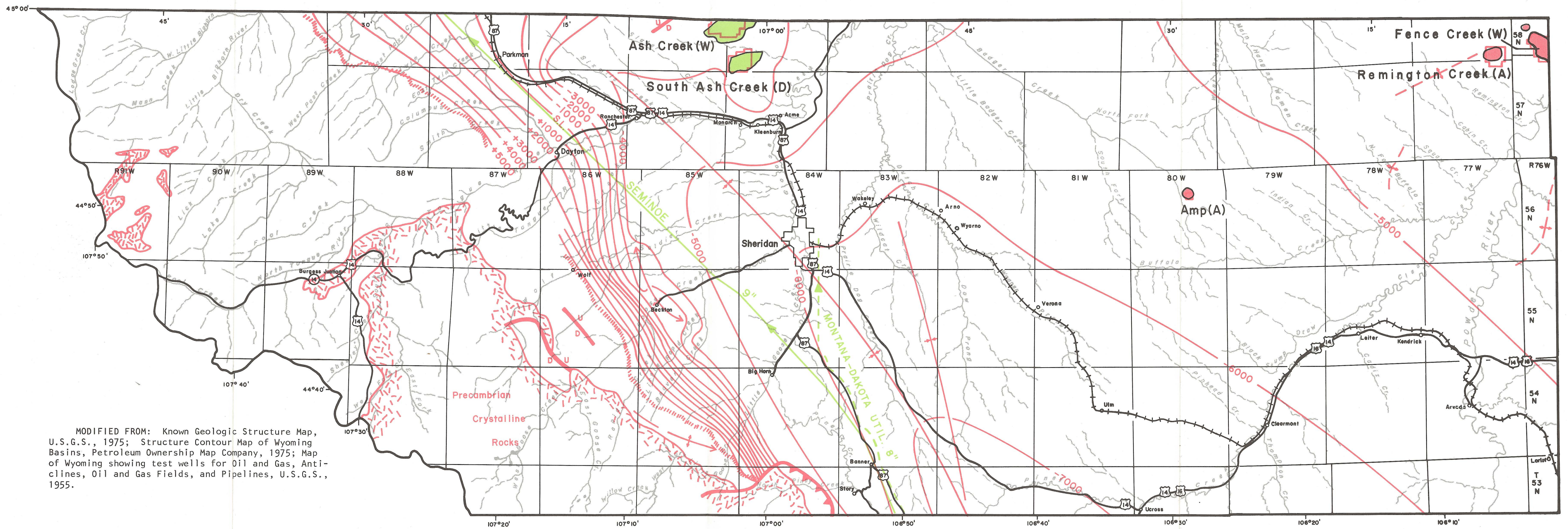
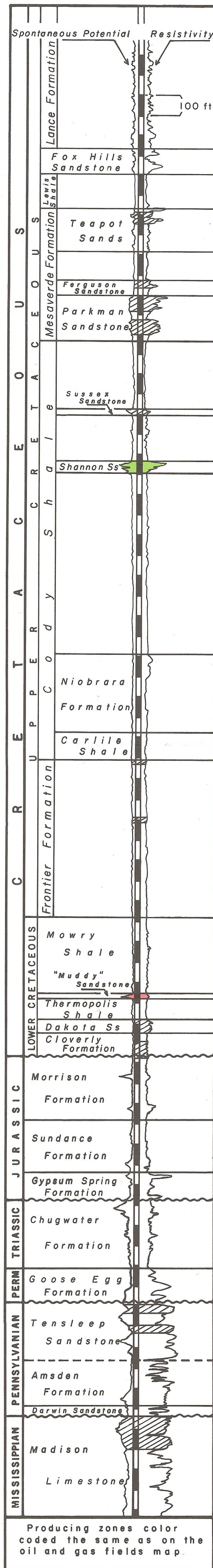
Location of known, tested, and potential sand and gravel deposits, as well as sources for crushable ledge rock, are shown on the map. The chart summarizes the characteristics of all the rock units found in Sheridan County.

Sources

Wyoming State Highway Department, 1965, Construction materials survey 1-90, Sheridan County, 185 p.

Wyoming State Highway Department, open file reports and maps.

Composite Electric Log



Oil and Gas Production

Field Name and Discovery Date	Location	Producing Formation	Average Producing Depth	No. Prod. Wells	Cumulative Production to Dec. 31, 1975 (BBL Gas/MCF)	Field Status
Amp (1970)	56N-80W	Muddy	9470'	0	980	545 Abnd.
Ash Creek (1952)	58N-84W, 85W	Shannon (Ash Creek)	4410'	7	5,384,722	0 Prod.(W)
Ash Creek (1954)	58N-84W	Shannon (Ash Creek)	4730'	8	7,993,306	0 Prod.(D)
Fence Cr. (1968)	58N-76W	Muddy	7710'	3	1,256,705	308,208 Prod.(W)
Remington Cr.	58N-76W	Muddy	8120'	0	72,051	8,917 Abnd.
Total of 5 Fields —				18	14,707,692	317,670

DATA SOURCE: State Board of Equalization, Ad Valorem Tax Department and the Wyoming Oil and Gas Conservation Commission.

PETROLEUM IN SHERIDAN COUNTY

Introduction

The first significant oil discovery in Sheridan County, made in 1952 approximately 15 miles north of Sheridan, was Ash Creek Field. It and South Ash Creek Field (discovered in 1954) are structural traps situated on a large, faulted anticlinal nose. Production is from the Upper Cretaceous Shannon Sandstone (locally referred to as the Ash Creek Sandstone) at an average depth of 4600 feet. Originally, 50 producing oil wells were drilled in the two fields, of which 15 are still producing.

Exploration continued without success through the 1950's and 1960's until 1968, when Fence Creek Field (partially in Campbell County) and Remington Creek Field were discovered in the northeastern part of the county. Both fields are stratigraphic traps producing from Lower Cretaceous Muddy Sandstone at average depths of 7710 and 8120 feet, respectively. Amp Field, a one well field discovered in 1970, produces from Muddy Sandstone at a depth of 9470 feet. These discoveries initially spurred active exploration for stratigraphic traps in the northern and eastern portions of the county. More recently, drilling activity has dropped off significantly and no new discoveries have been made since 1970.

Current oil production in the county is restricted to Ash Creek, South Ash Creek, and Fence Creek Fields. Of these, Ash Creek and Fence Creek Fields have active secondary (water flood) recovery programs underway. South Ash Creek Field, in addition to its oil production, serves as a water disposal project, allowing disposal of water produced with the oil in Ash Creek Field and South Ash Creek Field. Remington Creek and Amp Fields both produced for a short time, but are now abandoned.

As shown on the accompanying production graph, overall production has steadily declined except for the brief resurgence indicated for the late 1960's and early 1970's, resulting from the discovery of Fence Creek and Remington Creek Fields and the initiation of water flooding in Ash Creek Field. Recent initiation of water flooding in Fence Creek Field is probably responsible for the slight resurgence in production shown for 1976.

Sources of Production

Production to date within the county has been restricted to Cretaceous strata. The non-marine post-Cretaceous strata appear to lack favorable source rock and trapping mechanisms, and tests of pre-Cretaceous strata have thus far encountered either non-porous reservoir rock or poor source beds.

There are two types of traps present within the county, structural traps in the more highly deformed margin of the Powder River Basin and stratigraphic traps in the relatively stable interior portion of the Basin.

Cross-section B-B' illustrates the faulted anticlinal or structural trap forming Ash Creek and South Ash Creek Fields. Early exploration techniques centered around the search for surface expression of antidual structures, usually carried out by surface mapping and seismic work. The majority of structural traps in Wyoming have been discovered, and as a result newer, more sophisticated seismic techniques are now used to explore for the subtle, deep structural traps that remain to be found.

The stratigraphic trap, characteristic of the central Powder River Basin, is illustrated by cross-section A-A' through Fence Creek Field in the northeastern portion of Sheridan County. Most Muddy Sandstone reservoirs in the basin are stratigraphic traps consisting of thin beach sands or offshore bar sands that trend NW-SE, and thin discontinuous channel sands which are usually subperpendicular to beach and bar sands. The reservoir sandstones in Fence Creek Field are an example of NW-SE trending offshore bars that pinch out both upland and downland as shown on cross-section A-A'. Much of the present exploration in the Powder River Basin is for this type of stratigraphic trap.

Future Developments

When compared to oil and gas production to other, highly productive counties within the Powder River Basin, Sheridan County seems quite insignificant. Past drilling within the county has met with very limited success, and exploratory drilling has dropped off to the degree that, today, only one or two rigs are active at any given time. However, there are new discoveries in neighboring Campbell County that may be indicative of future success.

The best potential for future exploratory success probably lies in the search for new Cretaceous stratigraphic traps in the eastern and northern portions of the county. Judging from the numerous discoveries in adjacent northwestern Campbell County, the Muddy Sandstone still seems to hold potential. Improved seismic techniques could aid in exploration for deep structural traps along the western margin of the basin, on the flanks of the Bighorn Mountains.

Today, increasing oil and gas prices are fostering deep tests (greater than 20,000 feet) in other Wyoming basins, and it seems reasonable to expect deep drilling to become feasible in the Powder River Basin in the near future. In Sheridan County, areas near the axis of the basin and near the flanks of the Bighorn Mountains would probably be most affected by an increase in deep drilling, allowing for testing of potential Pennsylvanian, Tensleep and Mississippian Madison reservoirs. Potential accumulations in Lower Paleozoic pinch-outs and deep Cambrian Flathead reservoirs could also be tested. To date, the deepest drilling near the axis of the basin in Sheridan County has penetrated Lower Cretaceous strata.

Although the declining production trend could be reversed by one or two discoveries of the type just discussed, the county does not appear to have the potential for production that adjacent counties are exhibiting.

Selected References

Rocky Mountain Association of Geologists, 1972, Geologic atlas of the Rocky Mountain Region, 331 p.

Wyoming Geological Association, 1958, Thirteenth Annual Field Conference Guidebook: Powder River Basin, 341 p.

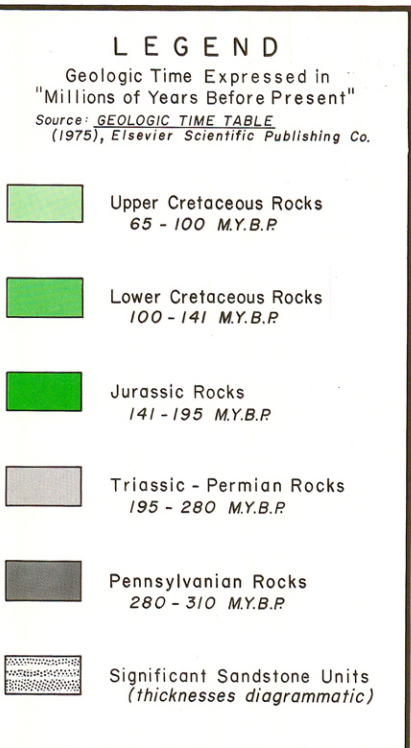
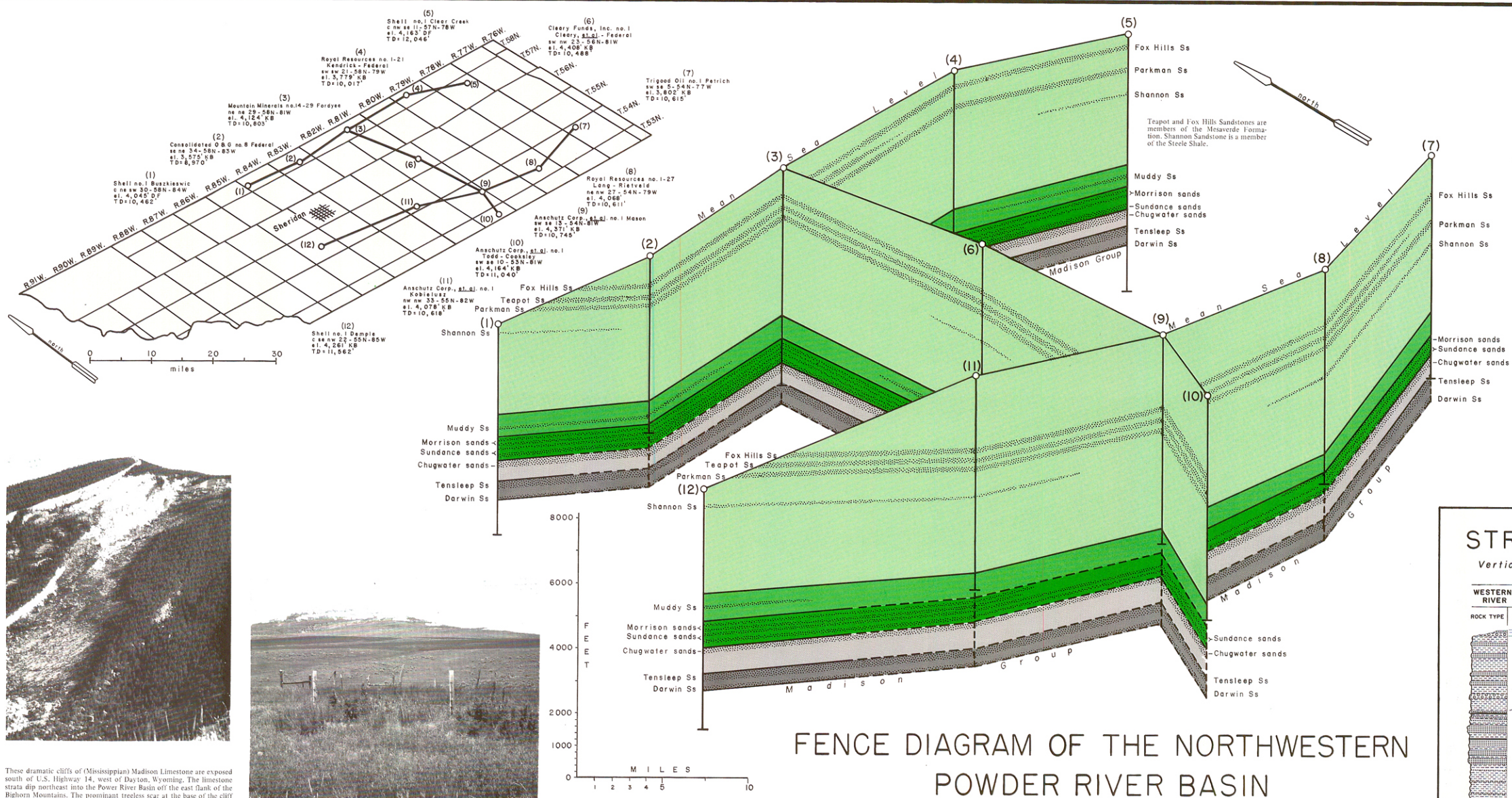
Wyoming Geological Association, 1962, Seventeenth Annual Field Conference Guidebook: Symposium on Early Cretaceous rocks of Wyoming, 339 p.

Wyoming Geological Association, 1976, Twenty-eighth Annual Field Conference Guidebook: Geology and energy resources of the Powder River Basin, 328 p.

Wyoming Geological Association and Billings Geological Society, 1963, First Joint Field Conference Guidebook: Northern Powder River Basin, Wyoming and Montana, 204 p.

Wyoming Oil and Gas Conservation Commission, 1977, Wyoming oil and gas statistics 1976, 125 p.

DNG



These dramatic cliffs of (Mississippian) Madison Limestone are exposed south of U.S. Highway 14, west of Dayton, Wyoming. The limestone strata dip northeast into the Powder River Basin off the east flank of the Bighorn Mountains. The prominent treeless scar at the base of the cliff is the result of landslide activity occurring along bedding planes.

Sandstones, siltstones, and claystones of the Wasatch Formation (Eocene) are exposed in this butte in eastern Sheridan County. Wasatch strata are nearly flat-lying, and cover most of Sheridan County east of the Bighorn Mountains. The formation is approximately 1700 feet thick and contains at least 18 coal beds of significant thickness (Culbertson and Mapel, 1976, p. 196). Thus, the Wasatch is a valuable resource for Sheridan County.

Photographs courtesy of David Gaylord, Remote Sensing Laboratory, University of Wyoming, Laramie.

FENCE DIAGRAM OF THE NORTHWESTERN POWDER RIVER BASIN (MEAN SEA LEVEL TO TOP OF MADISON GROUP)

STRATIGRAPHY OF SHERIDAN COUNTY, WYOMING

Sedimentary rocks, Cambrian to Recent in age, are exposed along the eastern and northern flanks of the Precambrian crystalline core of the Bighorn Mountains. These sedimentary rocks have been divided into 23 formations of 17,500 feet in aggregate thickness (Mapel, 1959, p. 9).

Core rocks are disconformably overlain by the Middle Cambrian Flathead Sandstone, 345 feet thick, a cross-bedded sandstone with interbedded micaceous shale and siltstone (Mapel, 1959, p. 10). The ratio of sand to shale may vary greatly between outcrops. The basal part of the Flathead is a pebbly sandstone, 10 to 40 feet thick.

The Flathead Sandstone is overlain by 645 feet of the Middle and Upper Cambrian Gros Ventre and Gallatin Formations. The Gros Ventre consists of thin-bedded calcareous, glauconitic sandstone, overlain by micaceous silty shales and thin-bedded, flat-pebble limestone conglomerate beds. The Gallatin is a cliff-forming, finely crystalline limestone containing pink, green, and red shale partings.

Cambrian strata are unconformably overlain by the Ordovician Bighorn Dolomite, up to 395 feet thick, a white, friable, cross-bedded sandstone overlain by 230 feet of massively bedded sugary dolomite characterized by a rough pitted weathered surface (Mapel, 1959, p. 13). Outcrops of the Bighorn generally form west-facing cliffs with east-facing gentle dip slopes.

The Bighorn Dolomite is overlain throughout most of Sheridan County by the Devonian Jefferson (or Darby) Formation, a limestone with thin interbedded shales and sandstones. Pinchout of Devonian strata in the Powder River Basin could result in petroleum entrapment.

Mississippian strata overlie the Bighorn in southern parts of the county, due to Late Devonian/Early Mississippian erosion and truncation. The Madison Limestone (Devonian and Mississippian) consists of carbonate rocks, up to 805 feet thick, in the northern Bighorn Mountains. The Madison has been divided into six members (Sando, 1976), on the basis of rock type and faunal content. The Madison Limestone is of current interest as an aquifer in the Powder River Basin.

The Madison Limestone is disconformably overlain by the Darwin Sandstone Member of the Amsden Formation at most localities, but locally by the Horseshoe Shale Member, or, rarely, the Ranchester Limestone Member, of the Amsden (Sando, 1976, p. 45). The disconformity between Madison and Amsden rocks is a karst surface formed on the Madison in Late Mississippian time. The Darwin is a cross-bedded conglomeratic sandstone, the lower part of which may be of fluvial origin. The Darwin is overlain by Amsden strata consisting of 240 feet of interbedded dolomite, red shale, finely crystalline limestone, calcareous sandstone, and siltstone (Mapel, 1959, p. 20).

Amsden strata are overlain by the Pennsylvanian Tensleep Sandstone, 275 feet thick, a well-sorted quartz sandstone. Large-scale cross-bedding is characteristic of the Tensleep. Pink to yellow dolomite with chert nodules is commonly interbedded with the sandstone.

Permian strata are overlain by 800 feet of the Triassic Chugwater Formation. The Chugwater is overlain by 800 feet of the Triassic Chugwater Formation. The Chugwater is overlain by 800 feet of the Triassic Chugwater Formation.

is 180 feet of claystone and shale interbedded with lenticular sandstone beds. Most sandstone beds are cross-bedded and contain chert pebbles, clay partings, and bone fragments (Mapel, 1959, p. 34).

Lower Cretaceous strata include the Cloverly Formation, Skull Creek Shale, Newcastle Sandstone, and Mowry Shale. The Cloverly consists of 150 feet of massive, cross-bedded sandstone, carbonaceous shale, and thin-bedded siltstone. Globular dalilite concretions are characteristic of the upper Cloverly. The Skull Creek Shale is 175 feet of nonresistant black shale interbedded with siltstone. The Newcastle Sandstone, 50 feet thick, is fine-grained sandstone with thin black shale partings. The Mowry Shale consists of a lower nonresistant black shale member, 200 feet thick, and an upper resistant light gray siliceous shale member, 325 feet thick.

Lower Cretaceous strata are overlain by the Upper Cretaceous Frontier Formation, Cody Shale, Parkman Sandstone member of the Mesaverde, Lewis Shale, and Lance Formation. The Frontier is 500 feet of interbedded shale and sandstone. A conglomeratic chert-pebble sandstone occurs at the top of the Frontier. The Cody Shale contains several members totaling 5570 feet in thickness (Mapel, 1959, p. 47), consisting of shale and sandstone beds. The Parkman Sandstone, 700 feet of massive sandstone overlain by gray shale and sandstone. The Lewis Shale, 700 feet thick, is a dark gray marine shale. The Lance Formation is 1960 feet of massive channelled sandstone and shale.

The Paleocene Fort Union Formation overlies Mesozoic strata. The Fort Union, 3900 feet thick, consists of 1500 feet of thin-bedded ferruginous sandstone, claystone, siltstone and fine-grained sandstone, overlain by 1500 feet of rock of much the same lithology with the addition of several thick carbonaceous shales, and 900 feet of poorly consolidated sandstone lenses, shale, and conglomerate (Mapel, 1959, p. 61).

The Eocene Wasatch Formation, which overlies the Fort Union unconformably, consists of sandstone, carbonaceous shale, and coal which interfingers to the west with conglomerate. The Kingsburg Conglomerate Member contains pebbles and cobbles of Paleozoic limestone and dolomite (Mapel, 1959, p. 63). The Kingsburg member represents coarse sediment eroded from the Rocky Mountain highlands during the Laramide Orogeny.

Oligocene and Miocene strata, the Arikaree, White River, and Moccasin Formations, represent a period of highland erosion and deposition in adjacent basins, which nearly buried the Rockies prior to late Cenozoic regional uplift. They consist mostly of conglomerate, sandstone, and volcanic ash.

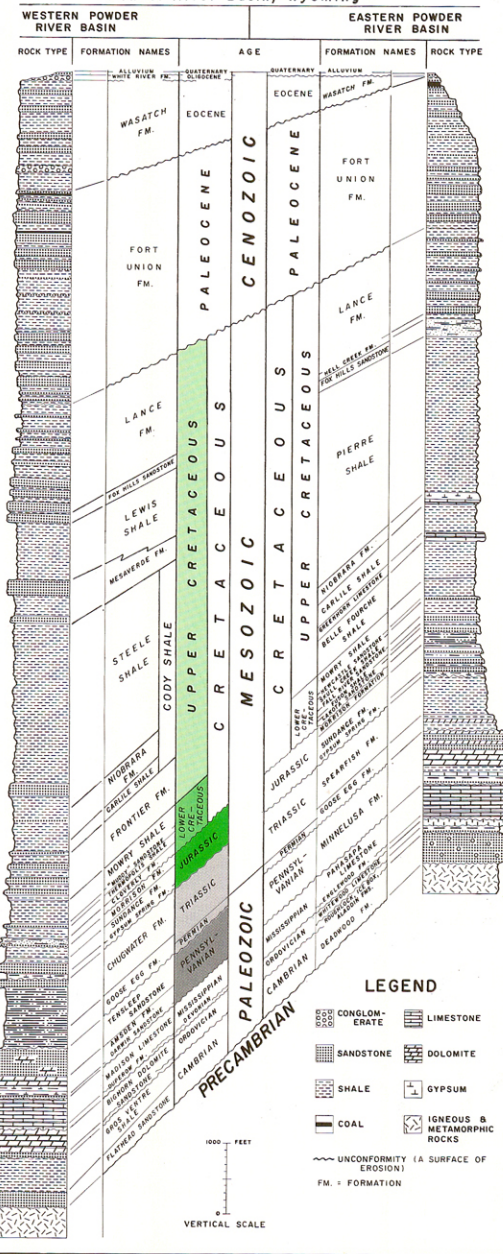
Pleistocene and Quaternary strata record episodes of glaciation, pedimentation and alluviation in, and adjacent to, the Bighorn Mountains.

REFERENCES

- Culbertson, W.C., and Mapel, W.J., 1976. Coal in the Wasatch Formation, northwestern part of the Powder River Basin near Sheridan, Sheridan County, Wyoming. Wyoming Geol. Assoc., 28th Annual Field Conference Guidebook, p. 193-201.
- Love, J.D., 1945. Stratigraphic sections and thickness maps of Triassic rocks in central Wyoming. U.S. Geol. Survey Oil and Gas Inv. (Prelim.) Chart 17.
- Mapel, W.J., 1959. Geology and coal resources of the Buffalo-Lake De Smet area, Johnson and Sheridan Counties, Wyoming. U.S. Geol. Survey Bull. 1075, 148 p., map.
- Sando, W.J., 1976. Madison Limestone (Devonian and Mississippian), east flank of Bighorn Mountains, Wyoming. Wyoming Geol. Assoc., 28th Annual Field Conference Guidebook, p. 45-52.
- Wyoming Geological Association, 1958, 13th Annual Field Conference Guidebook, Powder River Basin, 341 p.
- Wyoming Geological Association, 1976, 28th Annual Field Conference Guidebook, Geology and energy resources of the Powder River, 328 p.

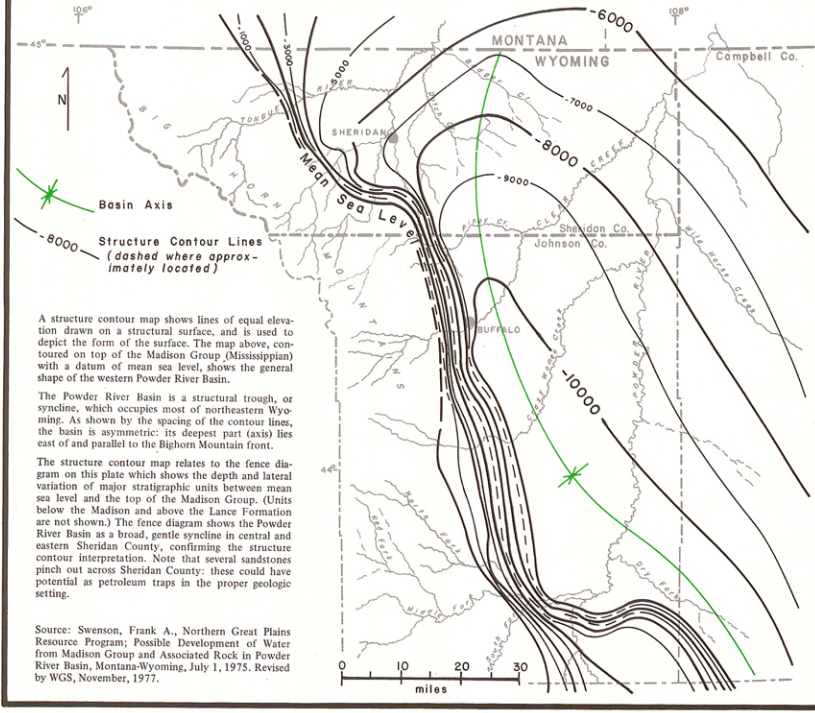
STRATIGRAPHIC COLUMN

Vertical Sequence Of Rock Formations Within The Powder River Basin, Wyoming



Structural Contour Map of the top of the Madison Group

(Contour Interval = 1,000 Feet)
Datum is Mean Sea Level



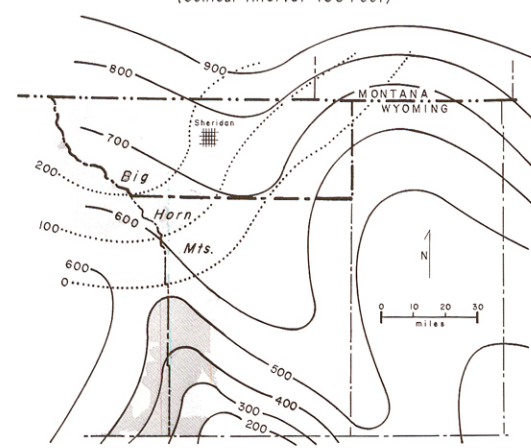
A structure contour map shows lines of equal elevation drawn on a structural surface, and is used to depict the form of the surface. The map above, contoured on top of the Madison Group (Mississippian) with a datum of mean sea level, shows the general shape of the western Powder River Basin.

The Powder River Basin is a structural trough, or syncline, which occupies most of northeastern Wyoming. As shown by the spacing of the contour lines, the basin is asymmetric: its deepest part (axis) lies east of and parallel to the Bighorn Mountain front.

Source: Swenson, Frank A., Northern Great Plains Resource Program; Possible Development of Water from Madison Group and Associated Rock in Powder River Basin, Montana-Wyoming, July 1, 1975. Revised by WGS, November, 1977.

Isopach Map of the Madison Group and Darby Formation

(Contour Interval = 100 Feet)



Isopach of the Madison Group (Mississippian).
Isopach of the Darby Formation (Devonian).
Source: Wyoming Geological Association Guidebook, 13th Annual Field Conference, 1958, p. 32 & 34. Reviewed by WGS, 1977.

Devonian rock units of northern Wyoming thin southward to a zero edge which trends across south-eastern Sheridan County, as shown by the dotted pattern on the isopach (thickness) map above. These rocks consist predominantly of marine carbonates which were deposited on a broad, shallow cratonic shelf bordering deeper waters to the west. In late Devonian or early Mississippian time, crustal movement caused uplift and erosion of parts of the cratonic shelf, and Devonian strata were completely eroded from central and southern Wyoming. Mississippian limestones were then deposited with regional unconformity on Devonian and older rocks across Wyoming. Subsequent periods of deformation have affected the regional thickness of Mississippian strata. Stratigraphic "pinch-outs," such as the southward thinning of Devonian strata, have potential as petroleum traps.

This map shows the major vegetation types in Sheridan County. The vegetation varies from the grasslands and widespread sagebrush-grasslands of the plains, to the alpine grasslands, meadows, and coniferous stands of the Bighorn Mountains. A classification scheme of twenty classes is employed, of which fourteen occur in Sheridan County. Vegetation types were distinguished by their appearance on aerial photographs (called their "visual aspect"). Density classes and principal plant species were derived from analysis of field data collected during the summer of 1976. Grass and herbaceous sites were sampled using the square-foot density method. Shrub and tree sites were sampled by recording the number of plants found within a sequence of randomly chosen square meter plots and 10 x 10 meter-square macroplots, respectively. Vegetation types and density classes are represented by two numbers, 7-3 for example, which refer to entries in the Vegetation Class and Density Class tables. Symbols for the principal plant species are four-letter codes representing the plants' scientific names, AGSM/BOGR for example, as listed in the table, Principal Plant Species of Sheridan County.

Explanation of Map Symbols Used For "Vegetation Type Designation" in Sheridan County, Wyoming.

SYMBOL ORDER

VEGETATION CLASS - DENSITY CLASS
PRINCIPAL PLANT SPECIES

DENSITY CLASSES**

CLASS	DEFINITION	GROUND COVER
1	trace/none	less than 5%
2	rare	5 - 25%
3	parklike	26 - 50%
4	interrupted	51 - 75%
5	continuous	76 - 100%

** Density classes refer to the percentage of ground covered by the total vegetation within a vegetation type.

VEGETATION CLASSES

- 1 Grassland and associated herbaceous plants, not under cultivation. Perennial grasses predominate and determine the aspect. Ex.: *Bouteloua gracilis*, *Agropyron smithii*
- 2 Alluvial areas; drainage from nearby highland (erosion usually heavy; sometimes saline).
- 3 Areas dominated by grasses rather than sedges. Generally moist, meadow-like areas in open timber or flood-plain areas, becoming moderately dry by mid-summer. Ex.: *Festuca idahoensis*, *Poa* spp.
- 4 Areas characterized principally by sedges, and by very poorly drained and/or partially submerged soils. Ex.: *Carex* spp.
- 5 Lakes, reservoirs, and other major water impoundments.
- * 6 Areas characterized by annual grasses or weeds. Also includes areas in transitory stages and semi-permanent conditions. Ex.: *Bromus tectorum*
- 7 Areas where sagebrush predominates.
- * 8 Areas where the various salt desert shrubs of the *Atriplex* family form the predominant vegetation and visual aspect. Ex.: *Atriplex nuttallii*
- * 9 Areas where greasewood (*Sarcobatus*) is the predominant vegetation and visual aspect. Includes valley floors subject to overflow during flood periods or areas underlain with ground water at shallow depths and having saline soil. Ex.: *Sarcobatus vermiculatus*
- 10 Areas where browse, except sagebrush or subtypes, gives the main aspect or is the predominant vegetation. Shrubs occupy the transition zone of lower mountain slopes, foothills, and plateau areas. Ex.: *Cercocarpus ledifolius*

- 11 Areas characterized by poplars (cottonwoods), willows, and other vegetation supported by stream water.
- 12 Areas characteristic of *Populus* spp. and other deciduous trees not including riparian.
- 13 Residential, business, and industrial areas (mines and mills included).
- 14 Areas where the land is cultivated for farmcrops, natural haylands, irrigated pastures, and fallow fields.
- 15 Areas characterized by coniferous species.
- * 16 Areas where timber has been logged.
- * 17 Areas of timberland, brushland or woodland that have been burned over, and a woody vegetation has not been established.
- * 18 Areas where vegetation is absent as a result of disturbance. Litter-covered ground is included.
- * 19 Areas of solid rock or land that is very rocky. Vegetation is essentially absent.
- 20 Areas where almost no vegetation exists and which have not been disturbed by man. Ex.: Intermittent lakebeds, rockslides, heavily eroded soils.

* NOTE: Vegetation classes, identified by asterisks (above), do not occur in Sheridan County singularly or as primary classes, but may occur as subordinate vegetation classes where other classes dominate.

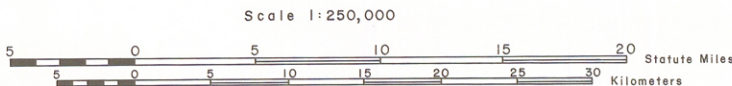
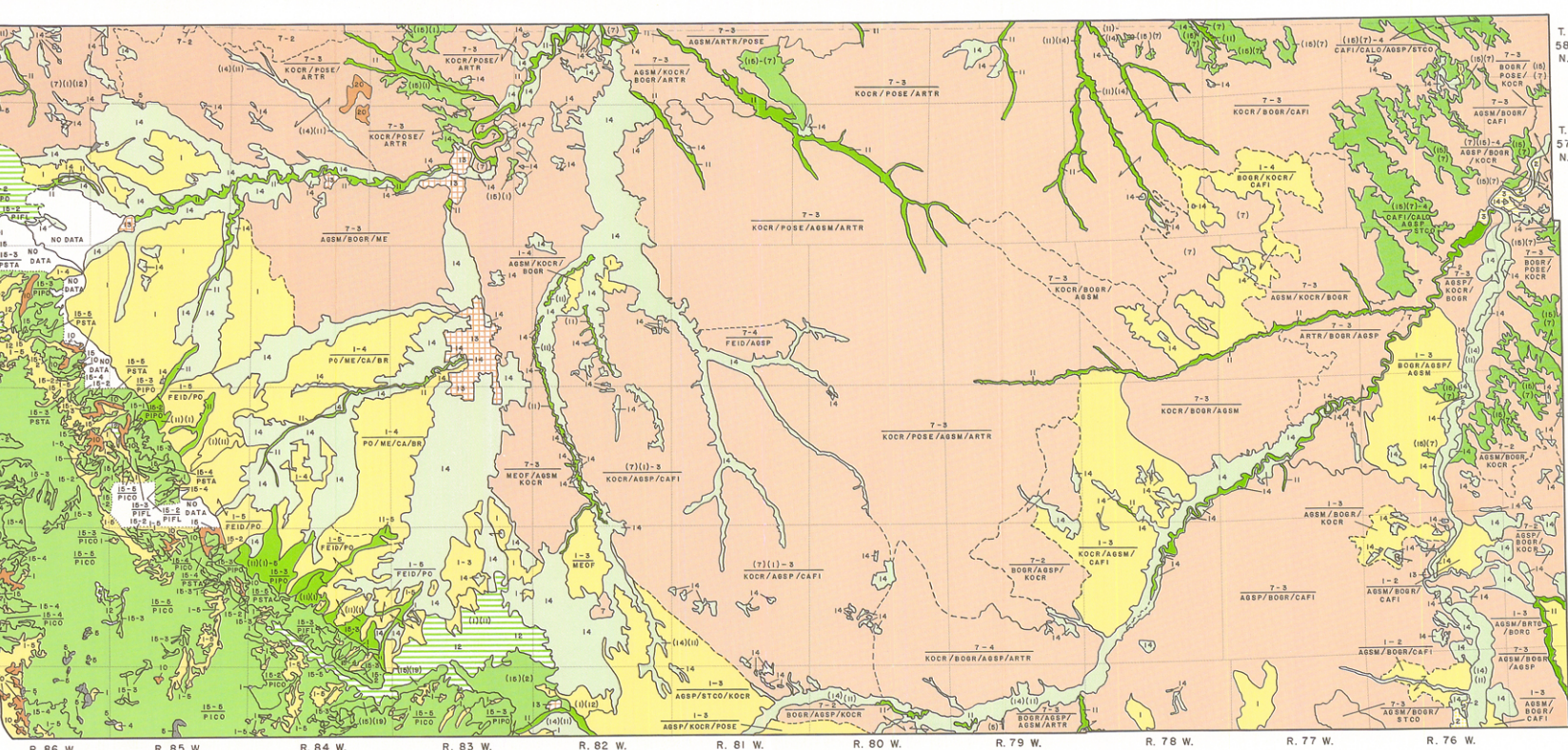
SOURCES OF DATA

- Beetle, A.A., 1970, Recommended plant names: Res. Journal No. 31, Agricultural Experiment Station, University of Wyoming.
- Inter-Agency Range Survey Comm., 1937, Instructions for range surveys.
- Kuchler, A.W., 1976, Vegetation mapping: The Ronald Press Company, New York.
- Thilenius, J.F., (Plant Ecologist), Personal communication, 1975: Rocky Mountain Forest and Range Experimental Station, Laramie, Wyoming.

Reliability Index



The vegetation map of Sheridan County was compiled from interpretations of high-altitude color infrared aerial photographs (NASA flights no. 73-147, August 30, 1973; and no. 310, May 14, 1975). Skylab color infrared photographs (S-190A, track 19, pass 6, June 6, 1973, frames 117-118) were used to cross check some areas in the northeastern portion of the county. Areas not adequately covered stereographically by high-altitude aerial photographs (cross hatched area on Reliability Index, above) were generalized from 1:24,000-scale U.S. Forest Service "forest-type" maps.



Vegetation map compiled by:
Alvin L. Medina
Department of Range Management
University of Wyoming

Work supported by:
U.S. Geological Survey
Grant 14-08-0001-G-163

PRINCIPAL PLANT SPECIES OF SHERIDAN COUNTY

SYMBOL	SCIENTIFIC NAME	COMMON NAME	SYMBOL	SCIENTIFIC NAME	COMMON NAME (con't.)
ABLA	<i>Abies lasiocarpa</i>	Alpine fir (subalpine f.)	HOBR	<i>Hordeum brachyantherum</i>	Meadow barley
ACNE	<i>Acer negundo</i>	Boxelder maple	HOJU	<i>Hordeum jubatum</i>	Foxtail barley
ACLA	<i>Achillea lanulosa</i>	Western yarrow	JU	<i>Juncus</i> (spp.)	Rush
ACHI	<i>Achillea millefolium</i>	Common yarrow	JUCO	<i>Juniperus communis</i>	Common juniper
AGGL	<i>Agoseris glauca</i>	Pale agoseris	JUSC	<i>Juniperus scopulorum</i>	Rockmountain juniper
AG	<i>Agropyron</i> (spp.)	Wheatgrass, quackgrass	KOCR	<i>Koeleria cristata</i>	Prairie junegrass
AGCR	<i>Agropyron cristatum</i>	Crested wheatgrass	LE	<i>Lepidium</i> (spp.)	Pepperweed
AGDA	<i>Agropyron dasystachyum</i>	Thickspike wheatgrass	LU	<i>Lupinus</i> (spp.)	Lupine
AGRE	<i>Agropyron repens</i>	Common quackgrass	MEBU	<i>Melica bulbosa</i>	Onion melic
AGSM	<i>Agropyron smithii</i>	Western wheatgrass	ME	<i>Melilotus</i> (spp.)	Sweetclover
AGSP	<i>Agropyron spicatum</i>	Bluebunch wheatgrass	MEOF	<i>Melilotus alba</i>	White sweetclover
AGAL	<i>Agrostis alba</i>	Redtop bent	MECI	<i>Melilotus officinalis</i>	Yellow sweetclover
ALCE	<i>Allium cernuum</i>	Nodding onion	OPPO	<i>Mertensia ciliata</i>	Mountain bluebells
AMAL	<i>Amelanchier alnifolia</i>	Saskatoon serviceberry	ORHY	<i>Opuntia polyacantha</i>	Plains pricklypear
ANGE	<i>Andropogon gerardi</i>	Big bluestem	PE	<i>Oryzopsis hymenoides</i>	Indian ricegrass
ANHA	<i>Andropogon hallii</i>	Sand bluestem	PHAL	<i>Penstemon</i> (spp.)	Penstemon
ANSC	<i>Andropogon scoparius</i>	Little bluestem	PH	<i>Phleum</i> (spp.)	Alpine timothy
ANRO	<i>Antennaria rosea</i>	Rose pussytoes	PIEN	<i>Phlox</i> (spp.)	Phlox
ARLO	<i>Aristida longiseta</i>	Red threawn	PICO	<i>Picea engelmanni</i>	Engelmann spruce
ARCO	<i>Arnica cordifolia</i>	Heartleaf arnica	PIFL	<i>Pinus contorta</i>	Lodgepole pine
ARCA	<i>Artemisia cana</i>	Silver sagebrush	PIPO	<i>Pinus flexilis</i>	Limber pine
ARFI	<i>Artemisia filifolia</i>	Sand sagebrush	PO	<i>Pinus ponderosa</i>	Ponderosa pine
ARFR	<i>Artemisia frigida</i>	Fringed sawwort	POCA	<i>Poa</i> (spp.)	Bluegrass
ARNO	<i>Artemisia nova</i>	Black sagebrush	POIN	<i>Poa canbyi</i>	Canby bluegrass
ARPE	<i>Artemisia pedatifida</i>	Birdfoot sawwort	POPR	<i>Poa interior</i>	Inland bluegrass
ARSP	<i>Artemisia spinescens</i>	Bud sawwort	POSE	<i>Poa pratensis</i>	Kentucky bluegrass
ARTR	<i>Artemisia tridentata</i>	Basin big sagebrush	POAN	<i>Poa secunda</i>	Sandberg bluegrass
AS	<i>Astragalus</i> (spp.)	Milkvetch	PODE	<i>Populus angustifolia</i>	Narrowleaf poplar
ATCO	<i>Atriplex confertifolia</i>	Shadscale saltbush	POTR	<i>Populus deltoides</i>	Eastern poplar
ATNU	<i>Atriplex nuttallii</i> (gardneri)	Muttail saltbush	POFR	<i>Populus tremuloides</i>	Quaking aspen
BASA	<i>Balsamorhiza sagittata</i>	Arrowleaf balsamroot	PR	<i>Potentilla fruticosa</i>	Shrubby cinquefoil
BEOC	<i>Betula occidentalis</i>	Water birch	PSME	<i>Prunus</i> (spp.)	Cherry, plum, etc.
BOCU	<i>Bouteloua curtipendula</i>	Sideoats grama	RO	<i>Pseudotsuga menziesii</i>	Common douglasfir
BOGR	<i>Bouteloua gracilis</i>	Blue grama	SA	<i>Quercus</i> (spp.)	Oak
BR	<i>Bromus</i> (spp.)	Brome	RO	<i>Rosa</i> (spp.)	Rose
BRAN	<i>Bromus anomalus</i>	Nodding brome	SA	<i>Salix</i> (spp.)	Willow
BRIN	<i>Bromus inermis</i>	Smooth brome	SATE	<i>Salsola tenuifolia</i>	Tumbleweed russiantistle
BRJA	<i>Bromus japonicus</i>	Japanese brome	SAVE	<i>Sarcobatus vermiculatus</i>	Black greasewood
BRMA	<i>Bromus marginatus</i>	Mountain brome	SE	<i>Sedum</i> (spp.)	Stonecrop
BRTE	<i>Bromus tectorum</i>	Cheatgrass brome	SHCA	<i>Shepherdia canadensis</i>	Russet buffaloberry
CACA	<i>Calamagrostis canadensis</i>	Bluejoint reedgrass	SIHY	<i>Sitonia hystrix</i>	Bottlebrush squirreltail
CAIN	<i>Calamagrostis inexpecta</i>	Northern reedgrass	SO	<i>Solidago</i> (spp.)	Goldenrod
CALO	<i>Calamovilfa longifolia</i>	Prairie sandreed	SOAM	<i>Sorbus americana</i>	American mountainash
CANU	<i>Calochortus nuttallii</i>	Sego mariposalily	STCO	<i>Stipa comata</i>	Needle and thread
CA	<i>Carex</i> (spp.)	Sedge	STVI	<i>Stipa viridula</i>	Green needlegrass
CAFI	<i>Carex filifolia</i>	Threadleaf sedge	SYAL	<i>Symphoricarpos albus</i>	Common snowberry
CAGE	<i>Carex geyeri</i>	Elk sedge	TAOF	<i>Taraxacum officinale</i>	Common dandelion
CAHE	<i>Carex heliophila</i>	Sun sedge	ZIVE	<i>Zigadenus venenosus</i>	Meadow deathcamus
CACO	<i>Castilleja coccinea</i>	Common indianpaintbrush			
CELE	<i>Cercocarpus ledifolius</i>	Curleaf mountainmahogany			
CEMO	<i>Cercocarpus montanus</i>	True mountainmahogany			
CHAL	<i>Chenopodium album</i>	Lambsquarters goosefoot			
CHLA	<i>Chrysothamnus lanceolatus</i>	Lanceleaf rabbitbrush			
CHNA	<i>Chrysothamnus nauseosus</i>	Rubber rabbitbrush			
CIAR	<i>Cirsium arvense</i>	Canada thistle			
CR	<i>Crataegus</i> (spp.)	Hawthorn			
DEBI	<i>Delphinium bicolor</i>	Little larkspur			
DISP	<i>Distichlis spicata stricta</i>	Inland saltgrass			
ELCA	<i>Elymus canadensis</i>	Canada wildrye			
ER	<i>Eriogonum</i> (spp.)	Wildbuckwheat			
EULA	<i>Eurotia lanata</i>	Common winterflat			
FEID	<i>Festuca idahoensis</i>	Idaho fescue			
GRSQ	<i>Grindelia squarrosa</i>	Curlycup gumweed			
GUSA	<i>Gutierrezia sarothrae</i>	Broom snakeweed			
HA	<i>Halogeton</i> (spp.)	Halogeton			

- Picture #1
Riparian site of the Powder River flood plain south of Arvada. The area is comprised of cottonwoods, willows, and assorted perennial and annual grasses and forbs (Class 11).
- Picture #2
Ponderosa pine--juniper--sagebrush association found on isolated breaks in the northeastern portion of the county and on the east side of the Powder River.
- Picture #3
Mountain scene showing four major vegetation classes (grassland, riparian, sagebrush-grassland, and coniferous), commonly found on the Bighorn Mountains (Classes 1, 11, 7, 15).



Picture # 1



Picture # 2



Picture # 3

